

Physically consistent responses in the atmospheric hydrological cycle in models and observations

Richard P. Allan

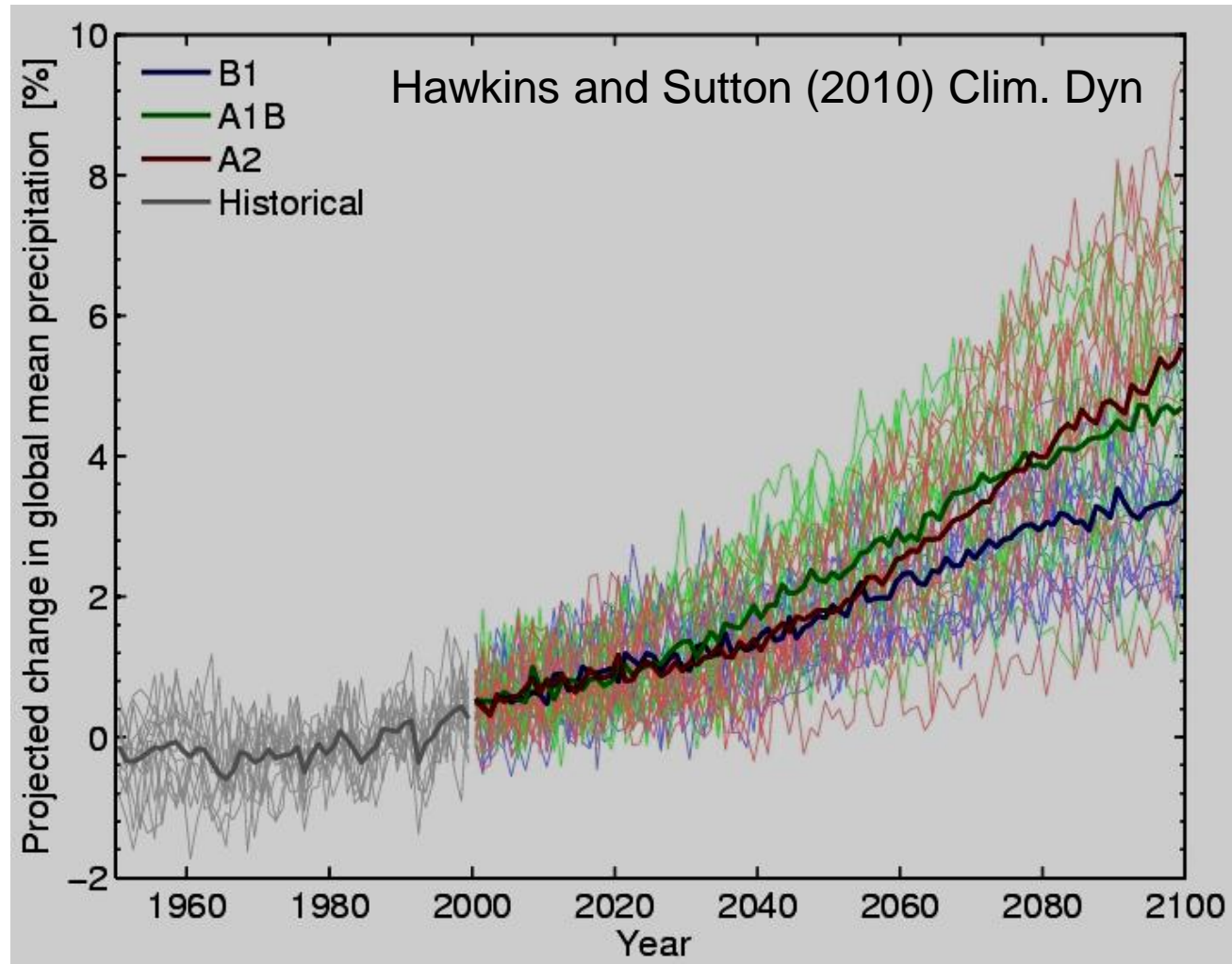
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<http://www.met.reading.ac.uk/~sgs02rpa>

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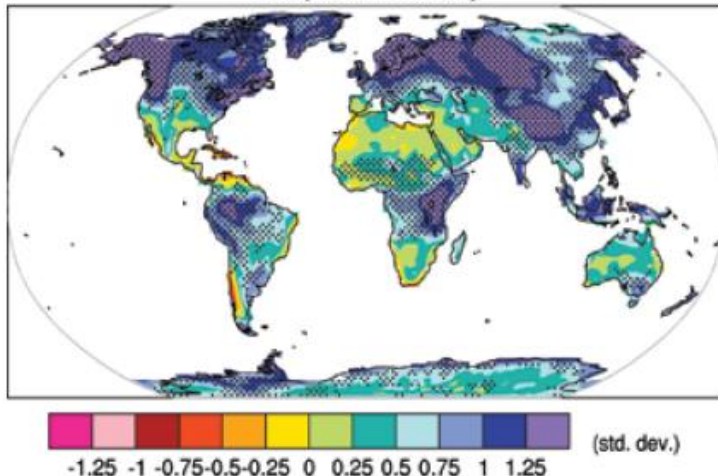
How should global precipitation respond to climate change?



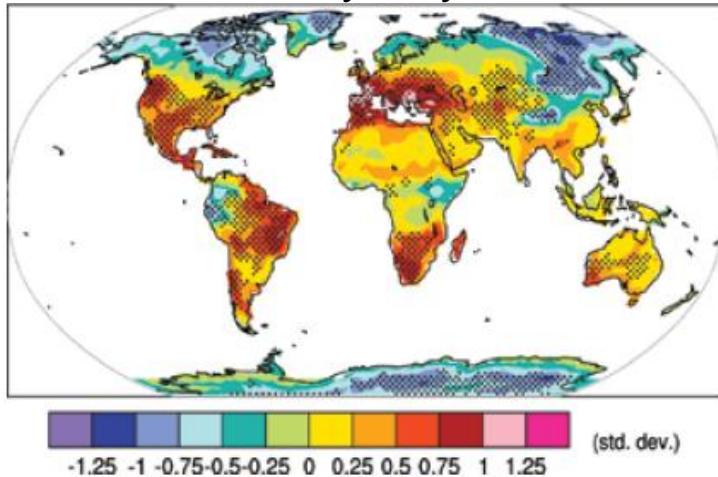
see also Allen and Ingram (2002) Nature; Trenberth (2011) Climate Research

Climate model projections (IPCC 2007)

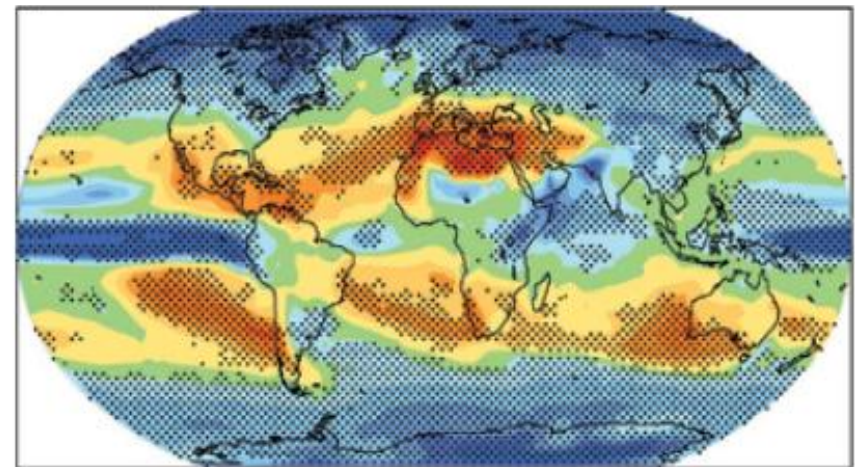
Precipitation Intensity



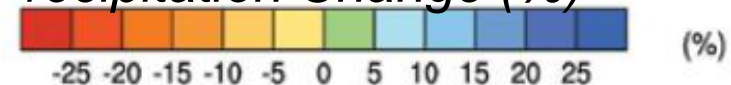
Dry Days



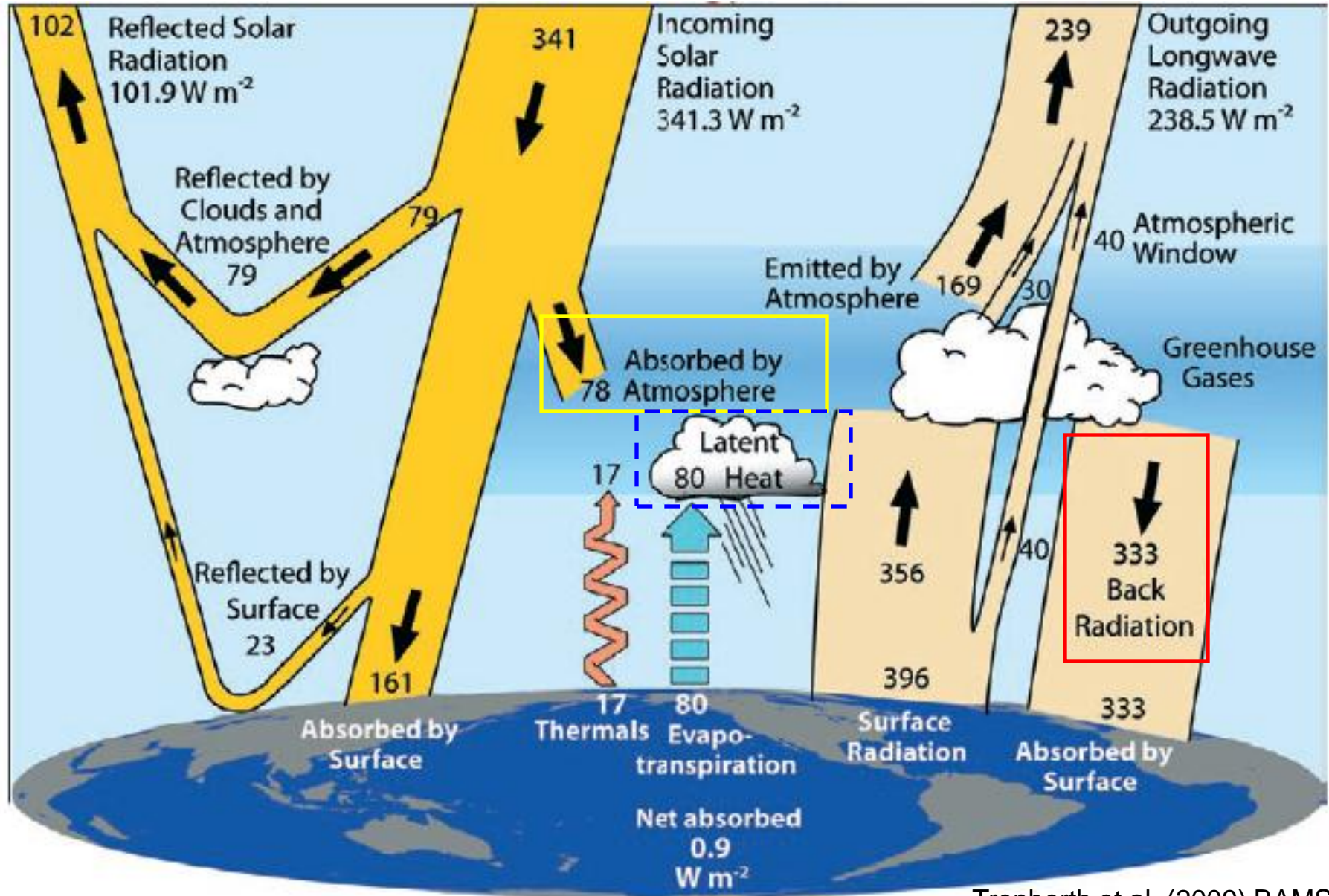
- Increased Precipitation
- More Intense Rainfall
- More droughts
- Wet regions get wetter, dry regions get drier?
- Regional projections??



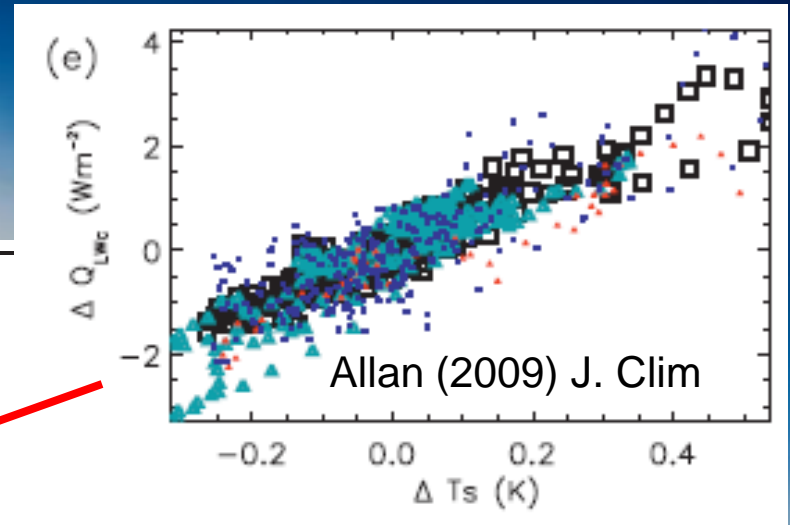
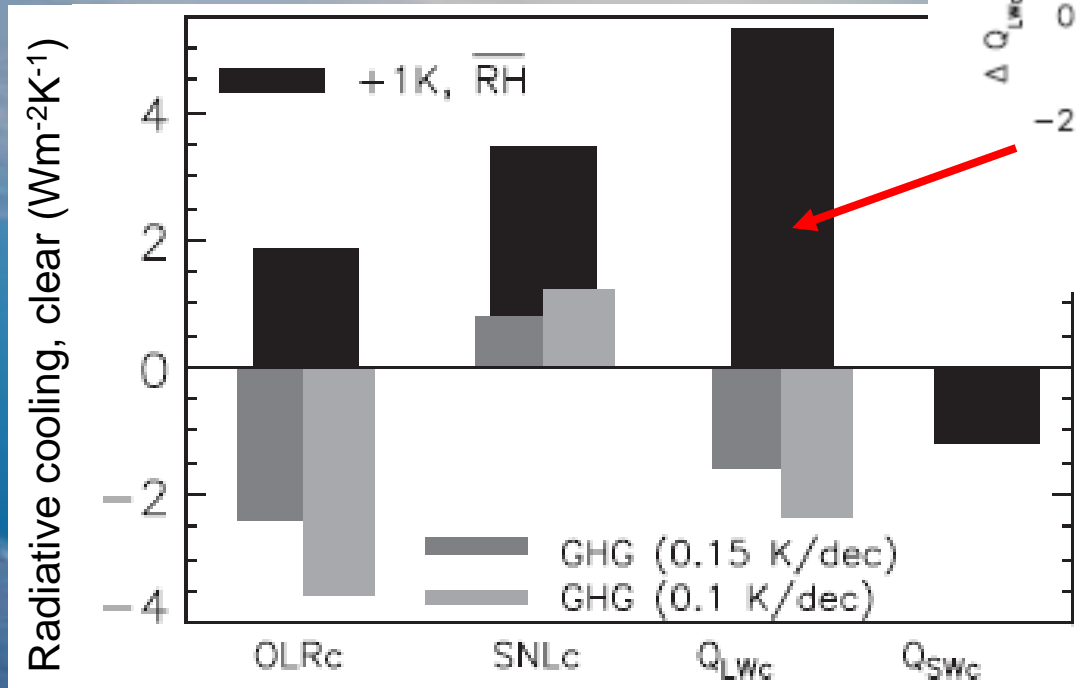
Precipitation Change (%)



Physical basis: energy balance



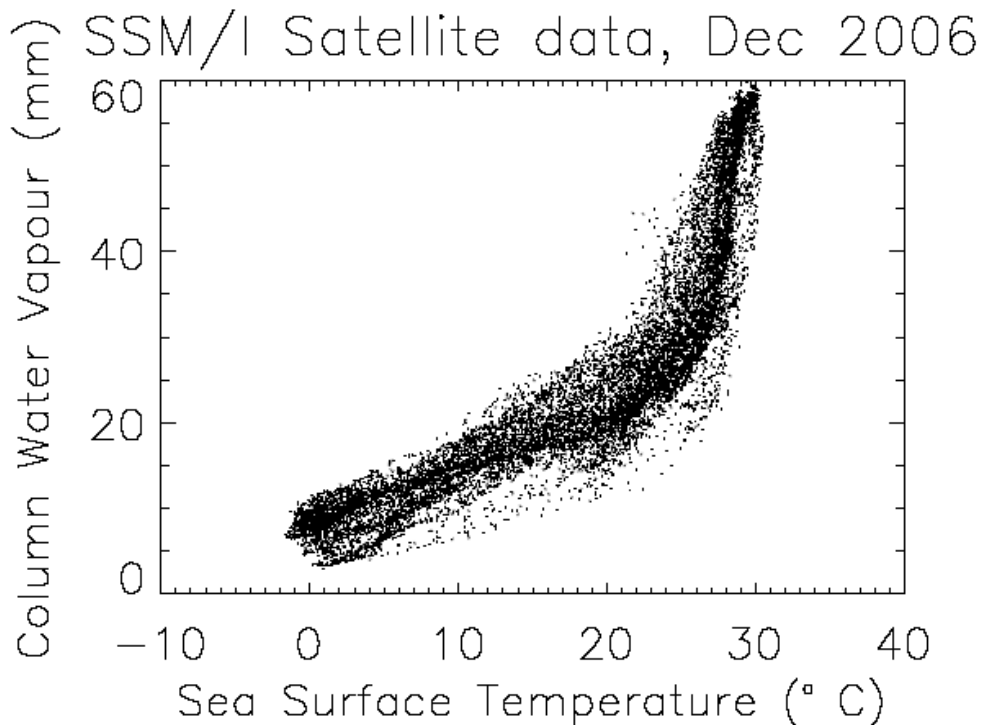
Models simulate robust response of clear-sky radiation to warming ($\sim 2 \text{ Wm}^{-2}\text{K}^{-1}$) and a resulting increase in precipitation to balance ($\sim 2 \text{ \%K}^{-1}$) e.g. Stephens & Ellis (2008) J Clim, Lambert and Webb (2008) GRL



$$\frac{dP}{dT_s} \sim \frac{1}{\rho_w L} \frac{dQ}{dT_s}$$

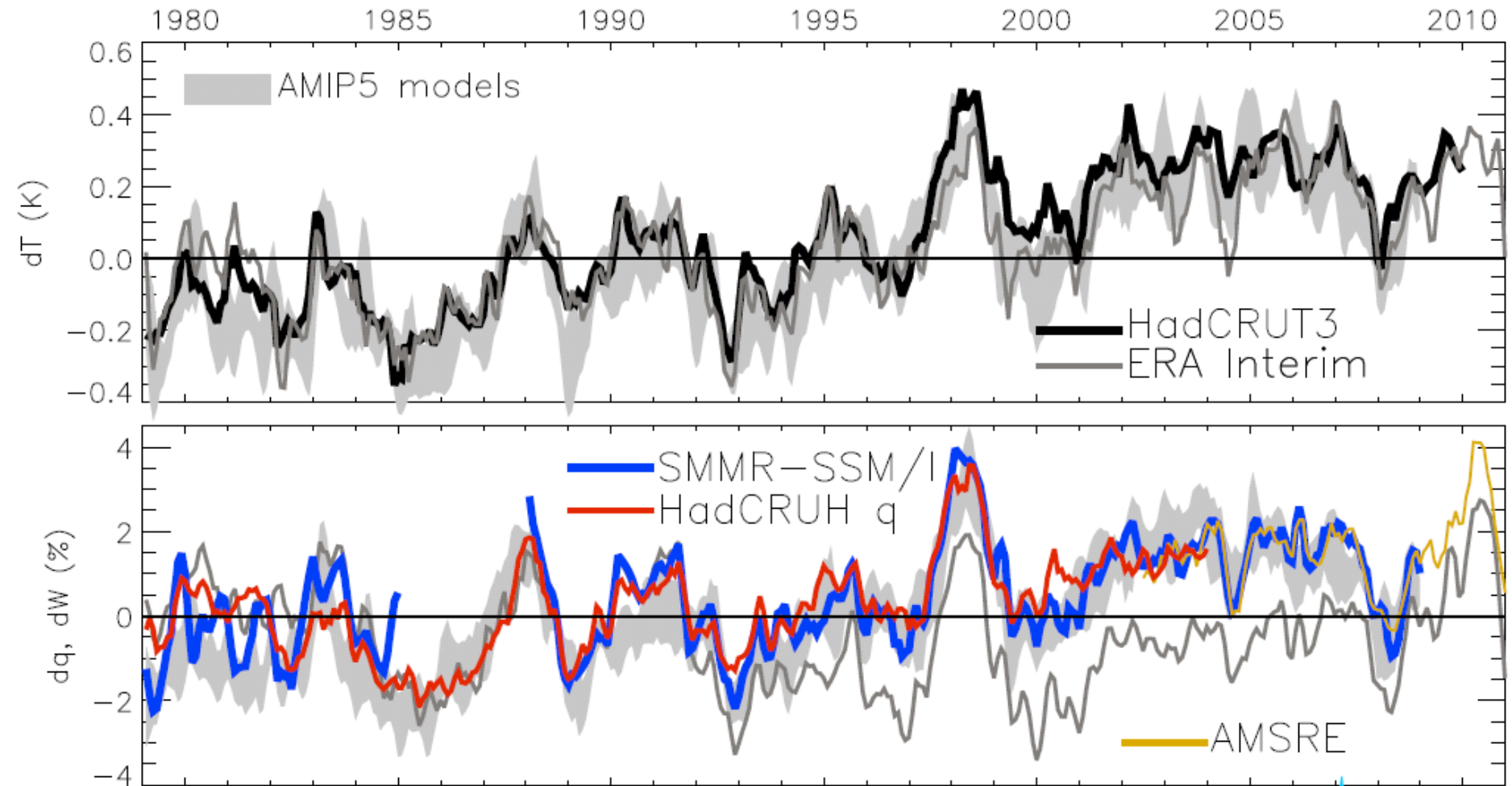
Physical basis: Clausius Clapeyron

$$\frac{1}{q_s} \frac{dq_s}{dT} \approx \frac{1}{e_s} \frac{de_s}{dT} = \frac{L}{R_v T^2} = \begin{cases} 0.14 K^{-1} & T = 200 K \\ 0.07 K^{-1} & T = 273 K \\ 0.06 K^{-1} & T = 300 K \end{cases}$$



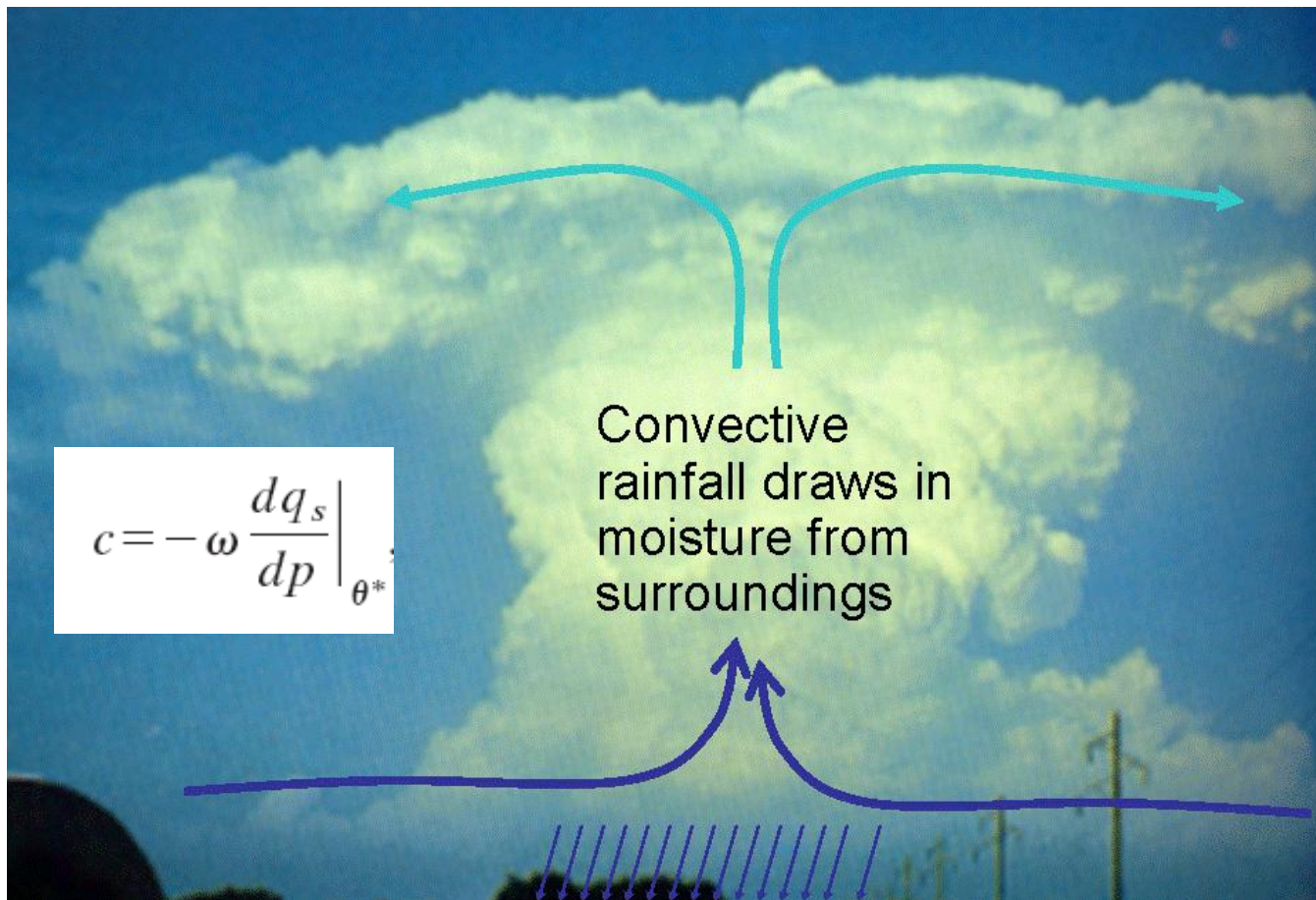
- Strong constraint upon low-altitude water vapour over the oceans
- Land regions?

Global changes in water vapour



Updated from [O'Gorman et al. \(2012\) Surv. Geophys](#); see also John et al. (2009) GRL

Extreme Precipitation



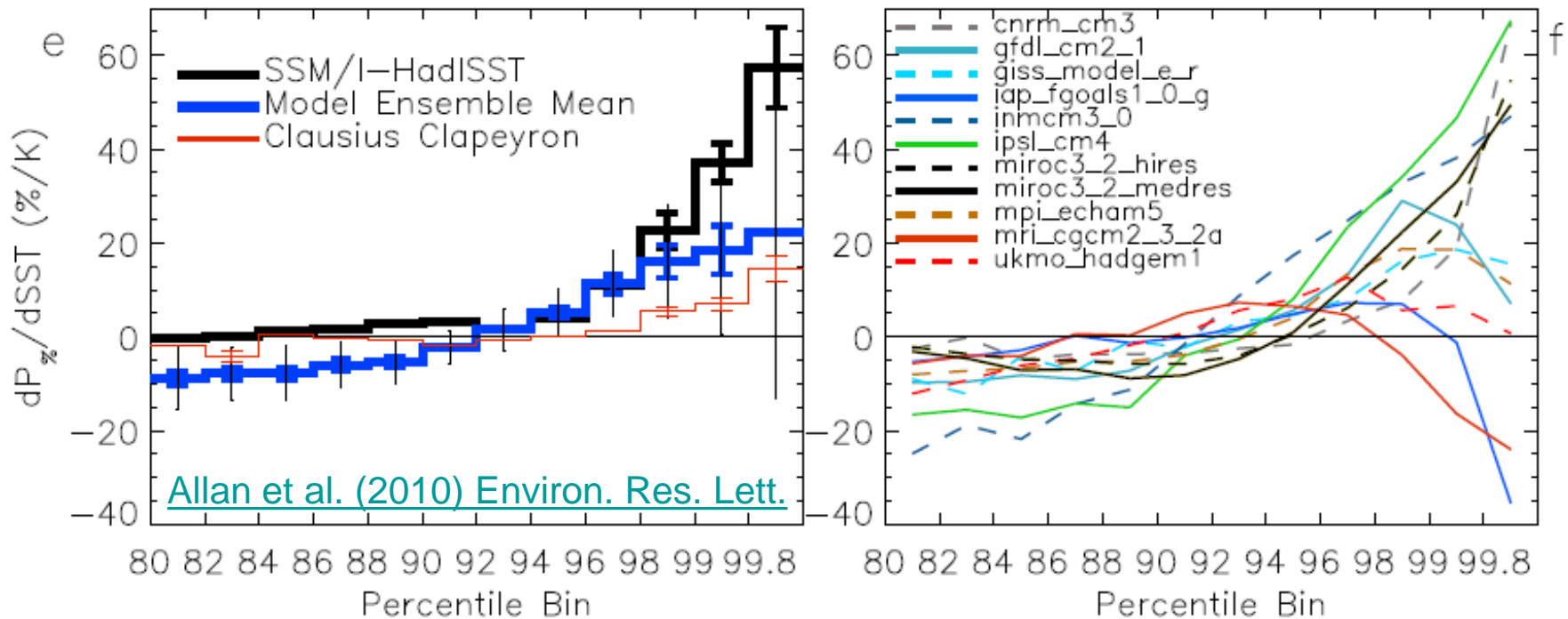
- Large-scale rainfall events fuelled by moisture convergence
 - e.g. [Trenberth et al. \(2003\) BAMS](#). But see [Wilson and Toumi \(2005\) GRL](#)

→ Intensification of rainfall ($\sim 7\%/K?$)

[O’Gorman and Schneider \(2009\) PNAS](#); [Gastineau and Soden \(2009\) GRL](#)

Observed and Simulated responses in extreme Precipitation

- Increase in intense rainfall with tropical ocean warming
- SSM/I satellite observations at upper range of models

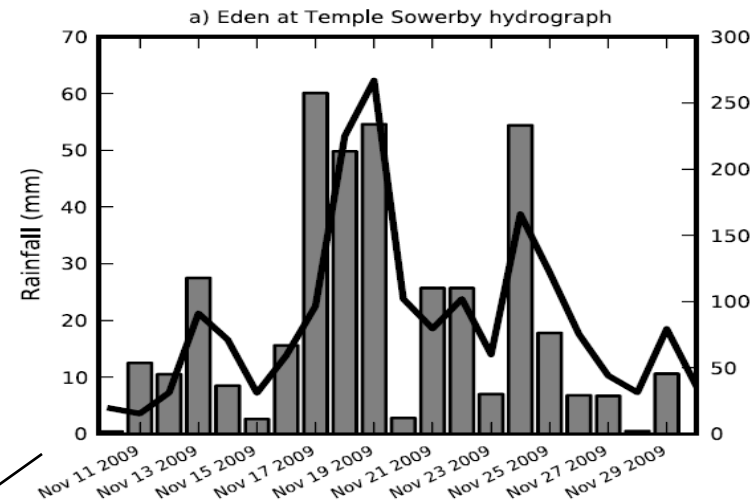


Tropical response uncertain: [O’Gorman and Schneider \(2009\) PNAS](#)....

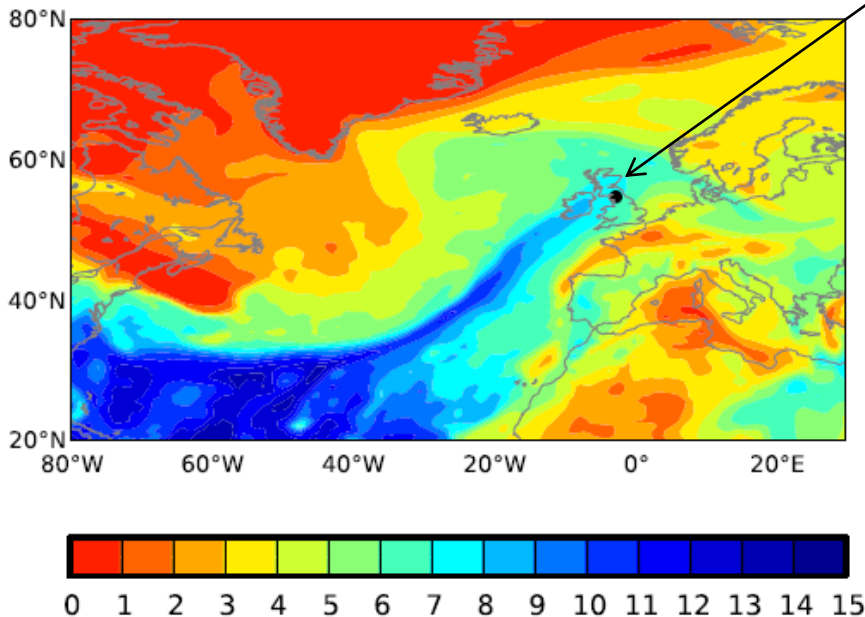
but see also: [Lenderink and Van Meijgaard \(2010\) ERL](#); [Haerter et al. \(2010\) GRL](#)

HydEF project: Extreme precipitation & mid-latitude Flooding

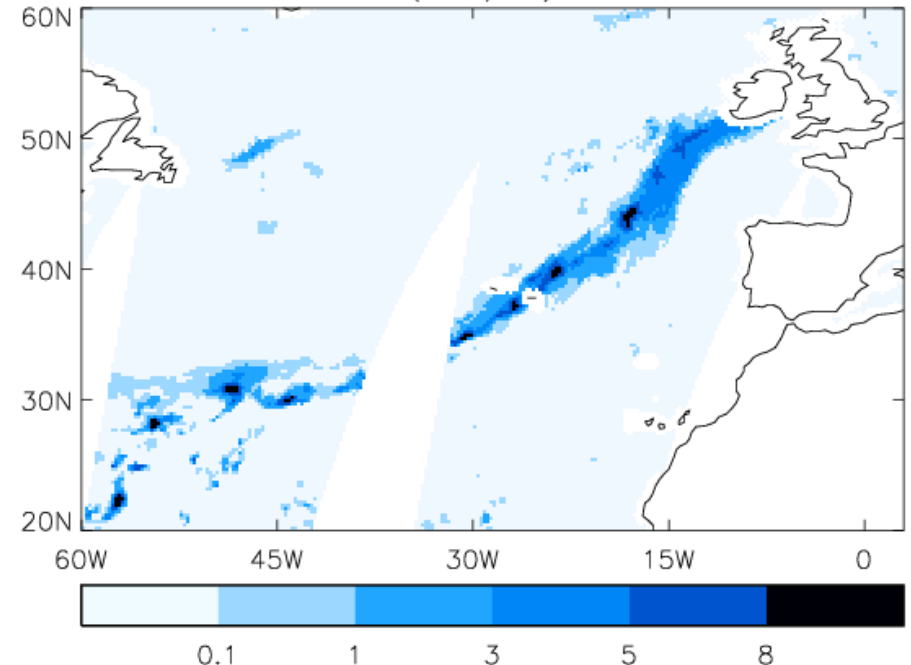
- Links UK winter flooding to moisture conveyor events
e.g. Nov 2009 Cumbria floods



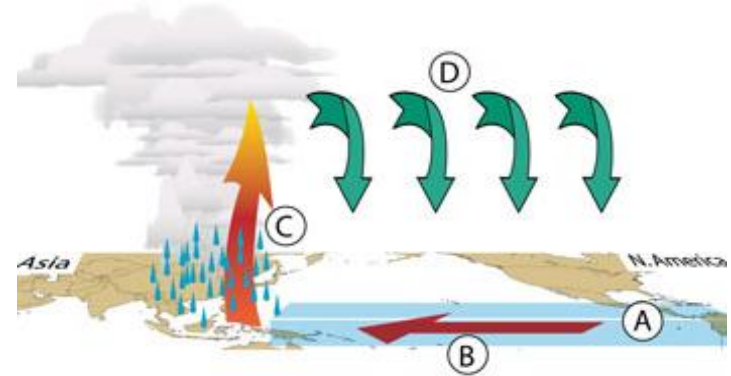
c) Specific humidity at 900 hPa (g kg^{-1})



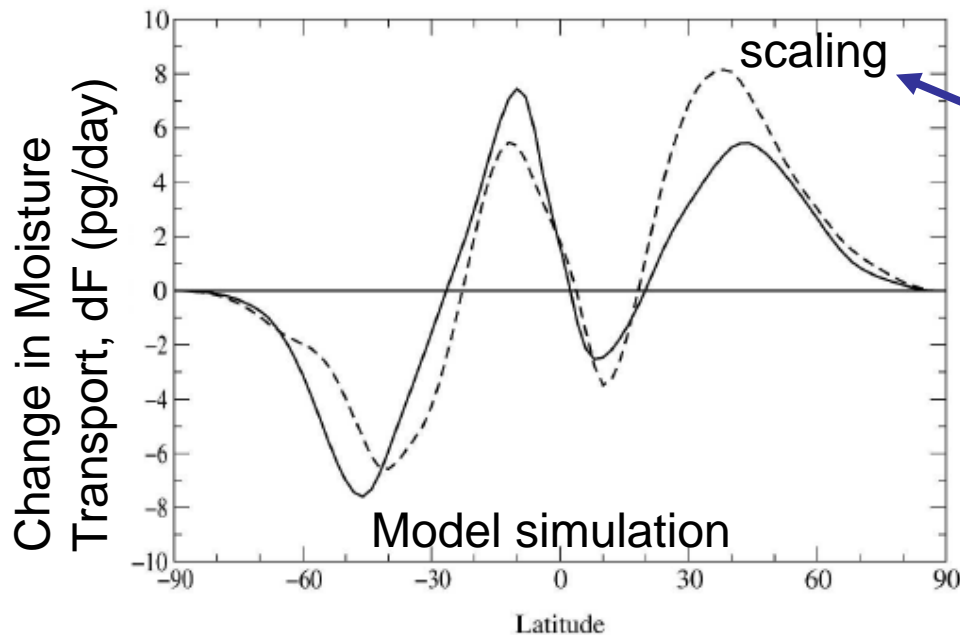
SSMIS F17 rainfall (mm/hr) 19 November 2009



Physical Basis: Moisture Balance



$$P-E \sim (\nabla \cdot (\mathbf{u} q)) \quad (\text{units of } s^{-1}; \text{ scale by } (p/g\rho_w) \text{ for units of mm/day})$$



$$\frac{\delta F}{F} \approx \frac{\delta e_s}{e_s} \approx \alpha \delta T.$$

$$\alpha \approx 0.07 \text{ K}^{-1}$$

If the flow field remains relatively constant, the moisture transport scales with low-level moisture.

Held and Soden (2006) J Climate

Projected (top) and estimated (bottom) changes in P-E

$$\frac{\delta F}{F} \approx \frac{\delta e_s}{e_s} \approx \alpha \delta T.$$

$$\delta(P - E) = -\nabla \cdot (\alpha \delta T F). \sim \alpha \delta T (P - E).$$

$$\alpha \approx 0.07 \text{ K}^{-1}$$

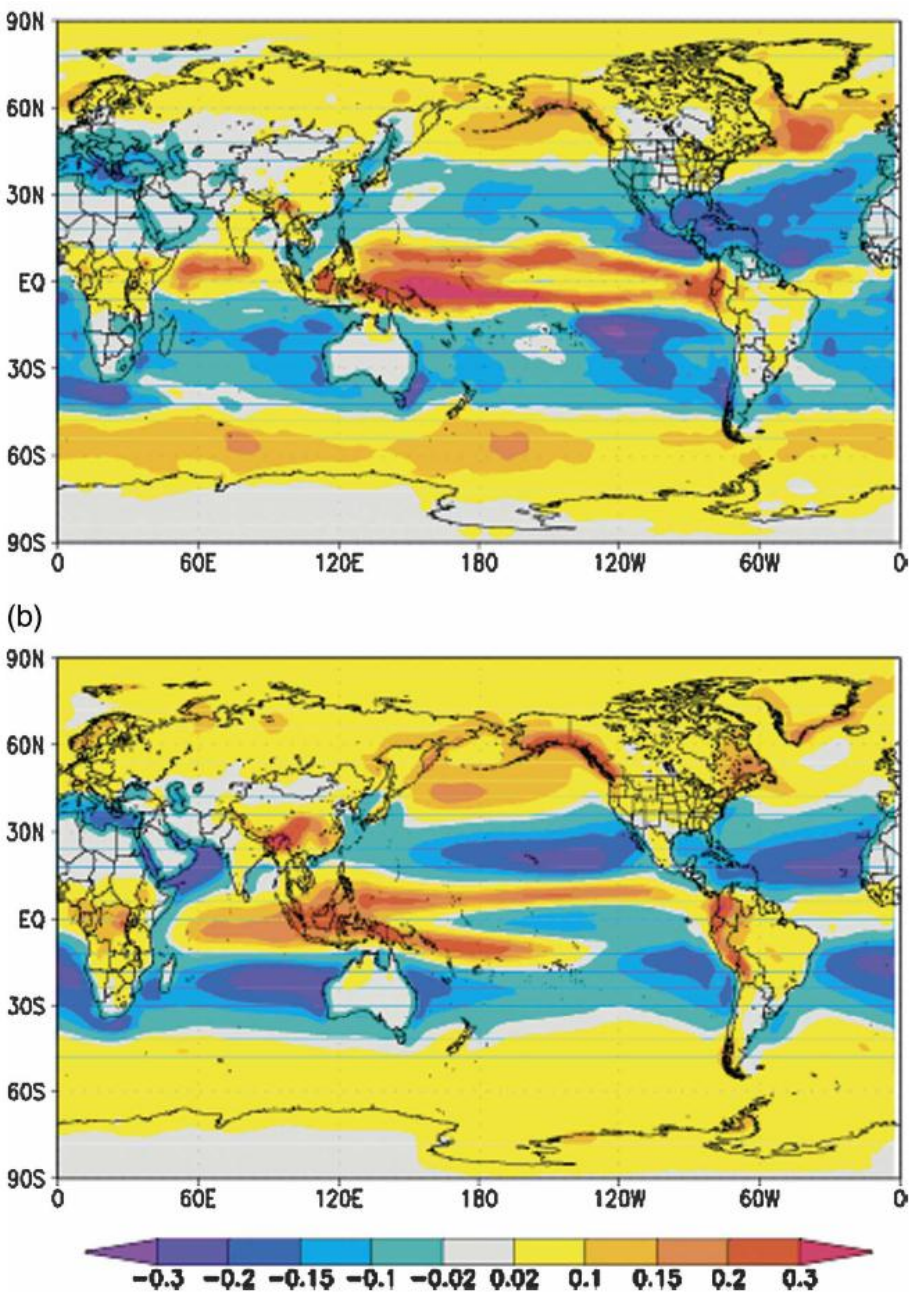
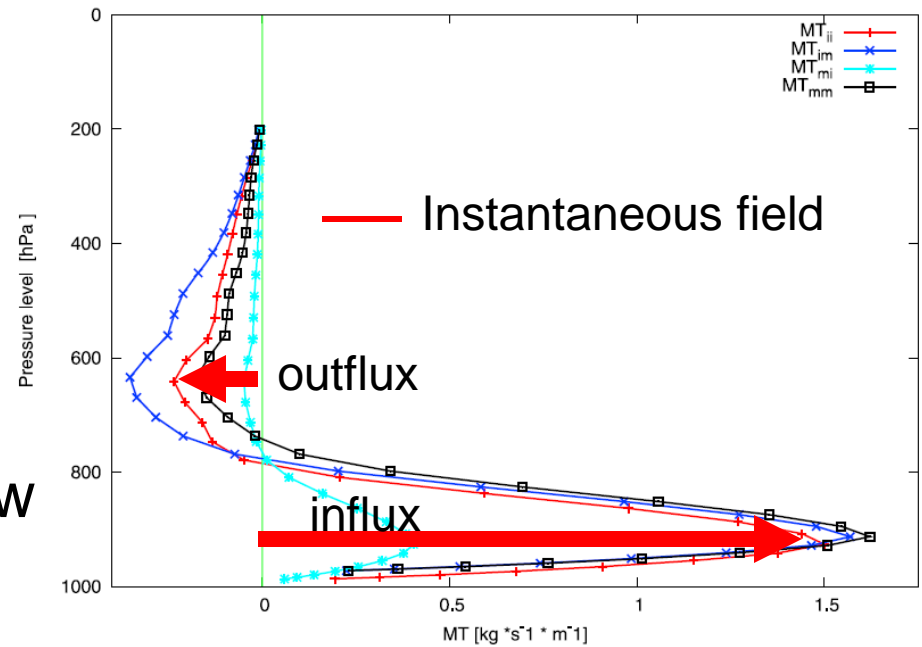


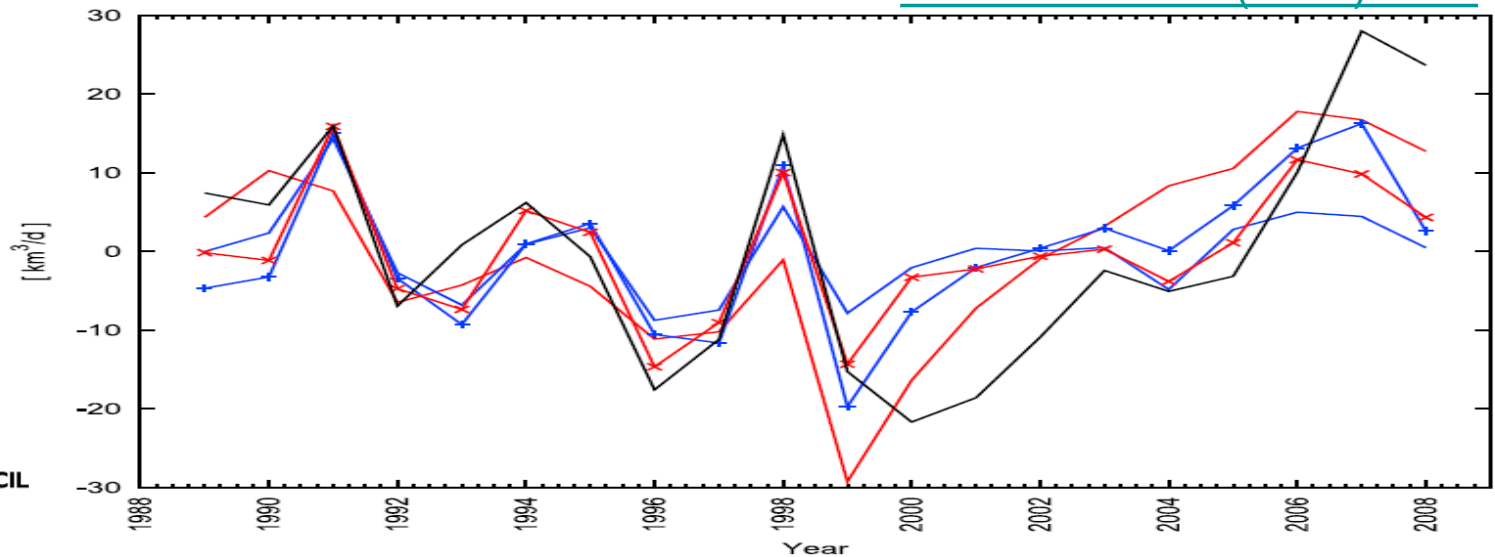
FIG. 7. The annual-mean distribution of $\delta(P - E)$ from the ensemble mean of (a) PCMDI AR4 models and (b) the thermodynamic component predicted from (6) from the SRES A1B scenario.

Moisture transports from ERA Interim

- Moisture transport into tropical ascent region
- Significant mid-level outflow
- 2000s: increases in inflow or drift in ERA Interim?



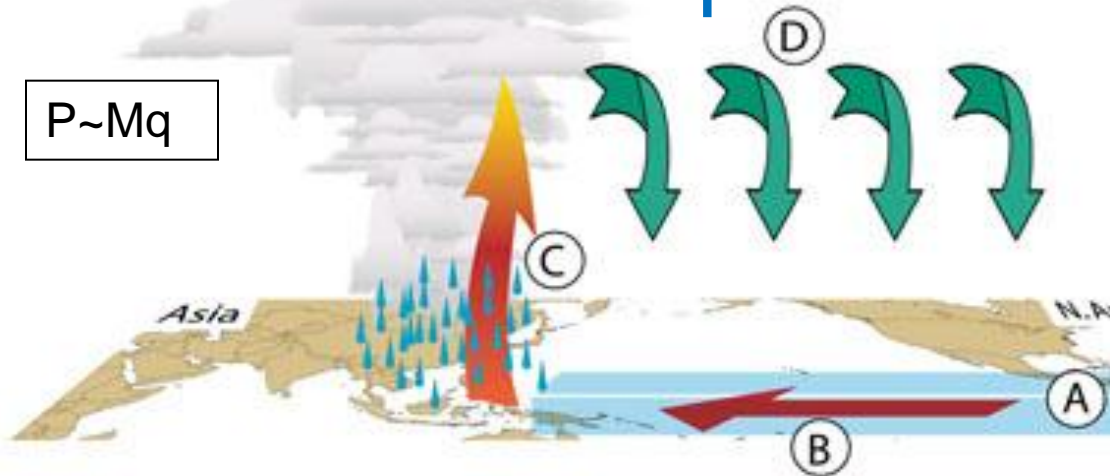
[Zahn and Allan \(2011\) JGR](#)



(a) yearly MT anomaly

Physical Basis: Circulation response

$$P \sim Mq$$



Walker circulation

- (A) Evaporation from warm ocean moistens lower atmosphere.
- (B) Trade winds carry moisture west
- (C) Moist air rises and feeds rain
- (D) Dry air cools and sinks

Warm climate

- (A) Atmospheric moisture increases strongly.
- (C) Rainfall increases more slowly than moisture

To compensate, winds slow.

First argument:

$$P \sim Mq.$$

So if P constrained to rise more slowly than q , this implies reduced M

Second argument:

$$\omega = Q/\sigma.$$

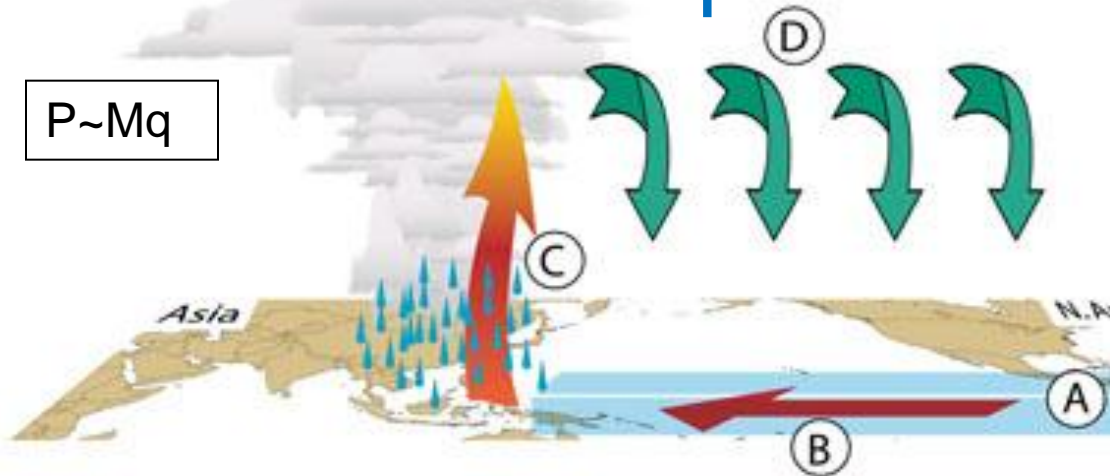
Subsidence (ω) induced by radiative cooling (Q) but the magnitude of ω depends on $(\Gamma_d - \Gamma)$ or static stability (σ).

If Γ follows MALR \rightarrow increased σ . This offsets Q effect on ω .

See Held & Soden (2006) and Zelinka & Hartmann (2010) JGR

Physical Basis: Circulation response

P~Mq



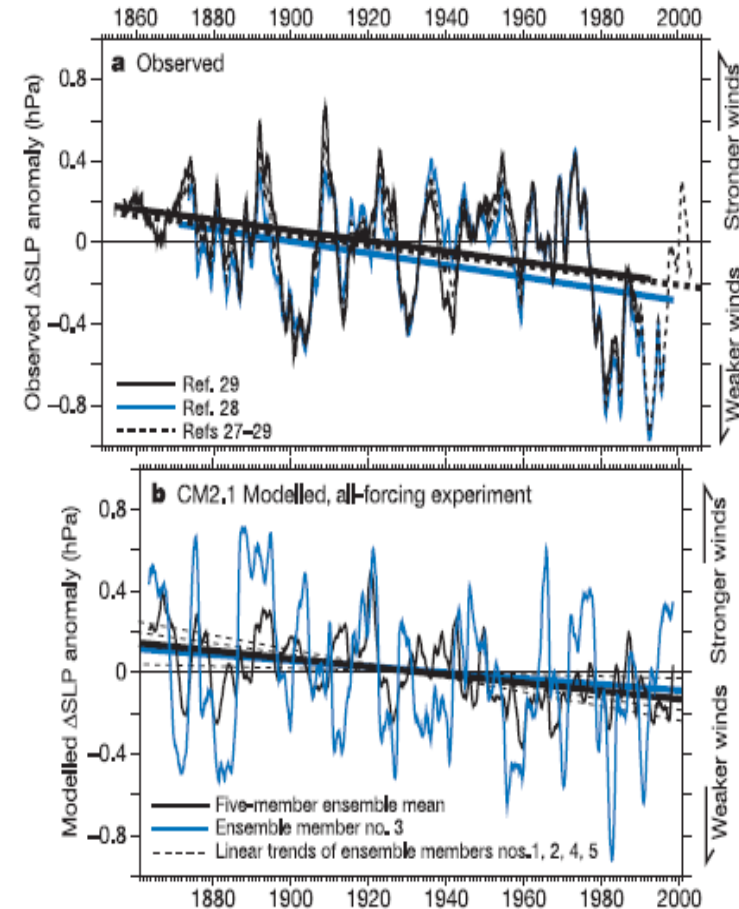
Walker circulation

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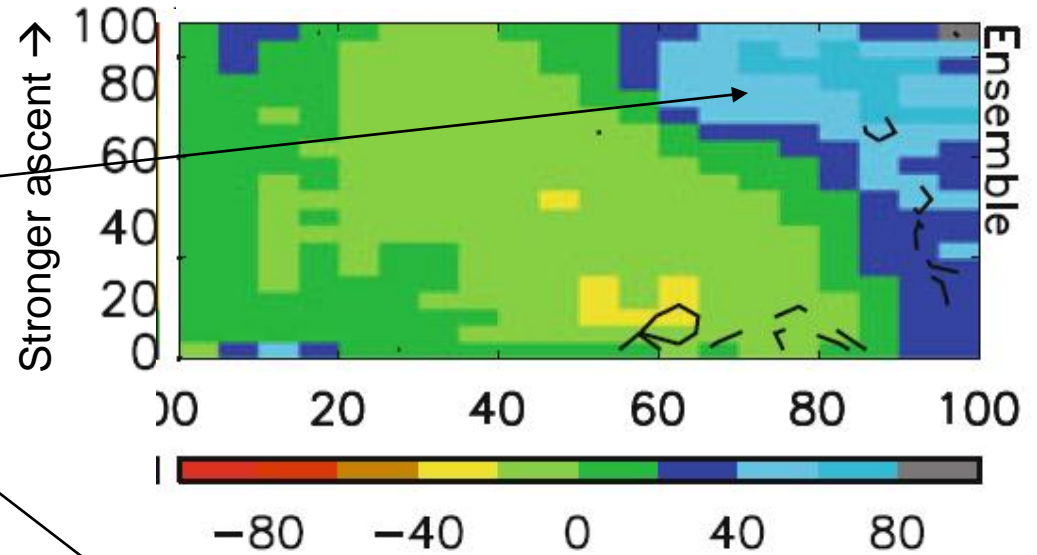
Models/observations achieve muted precipitation response by reducing strength of Walker circulation. Vecchi and Soden (2006) Nature

Precipitation bias and response binned by dynamical regime

- Model biases in warm, dry regime
- Strong wet/dry fingerprint in model projections (below)

[Allan \(2012\) Clim. Dyn.](#)

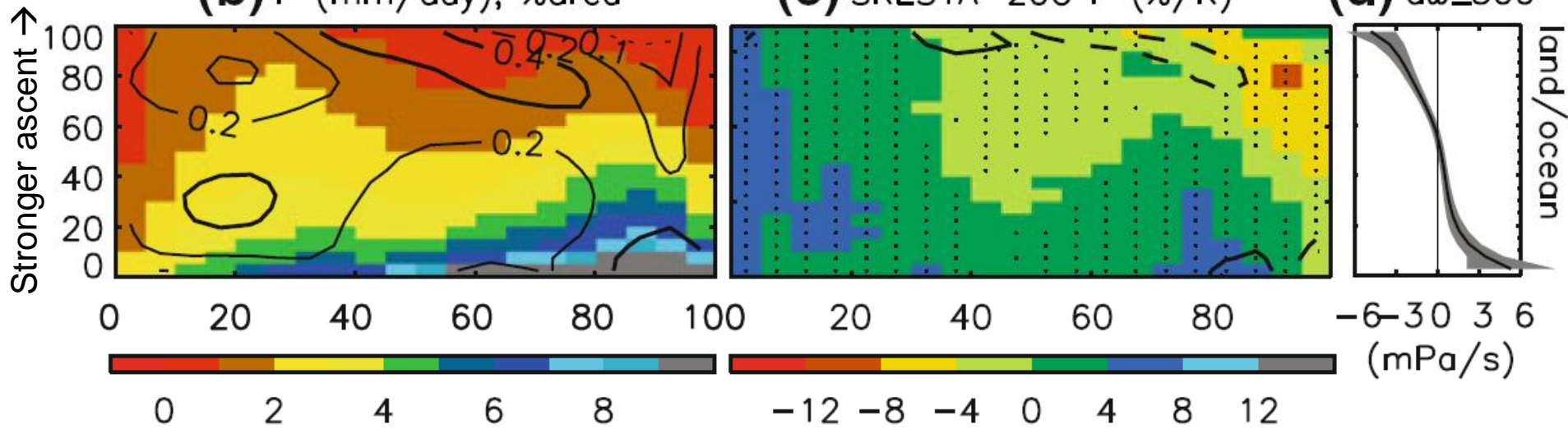
Warmer surface temperature →



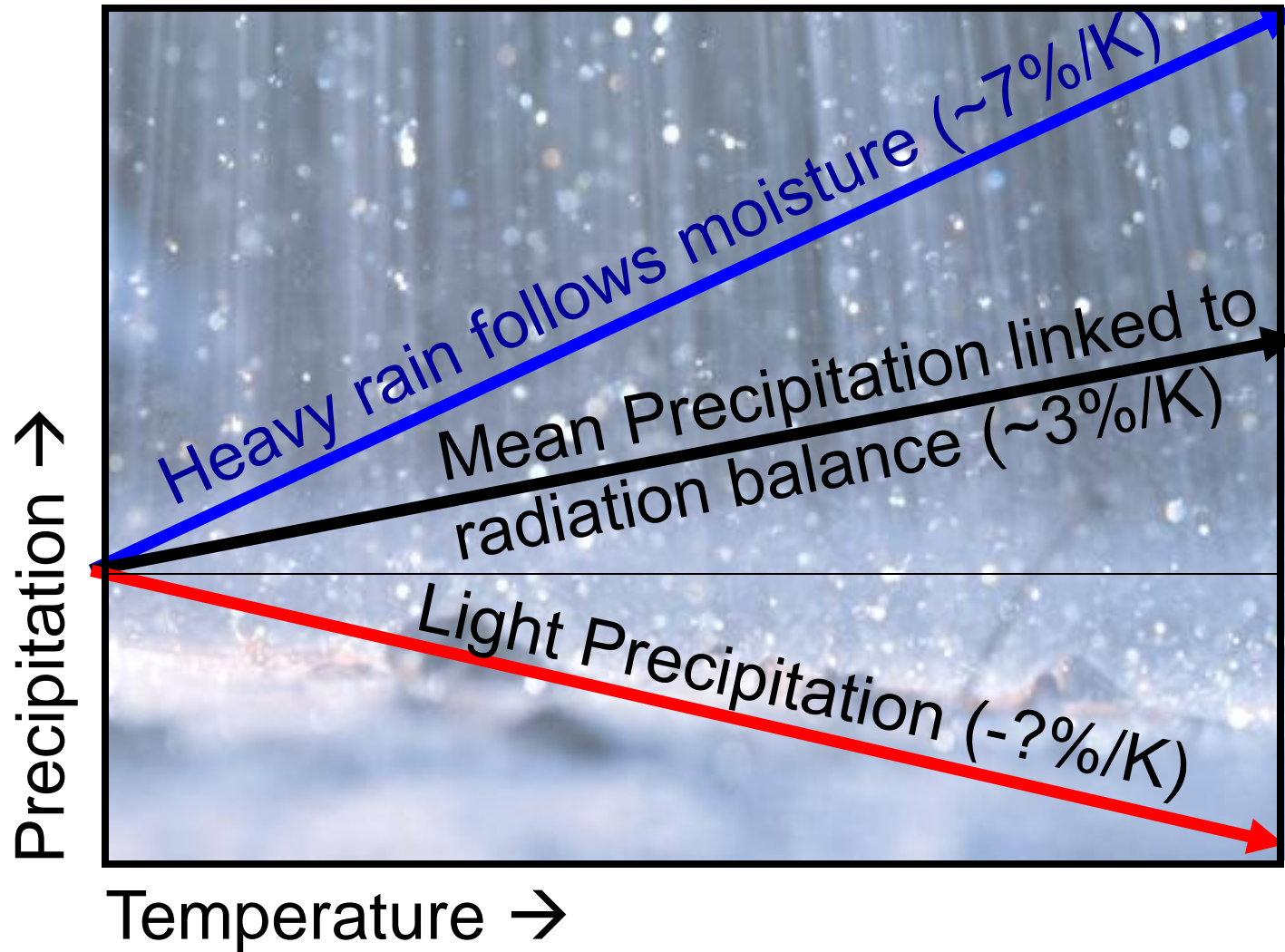
(b) P (mm/day); %area

(c) SRES1A-20C P (%/K)

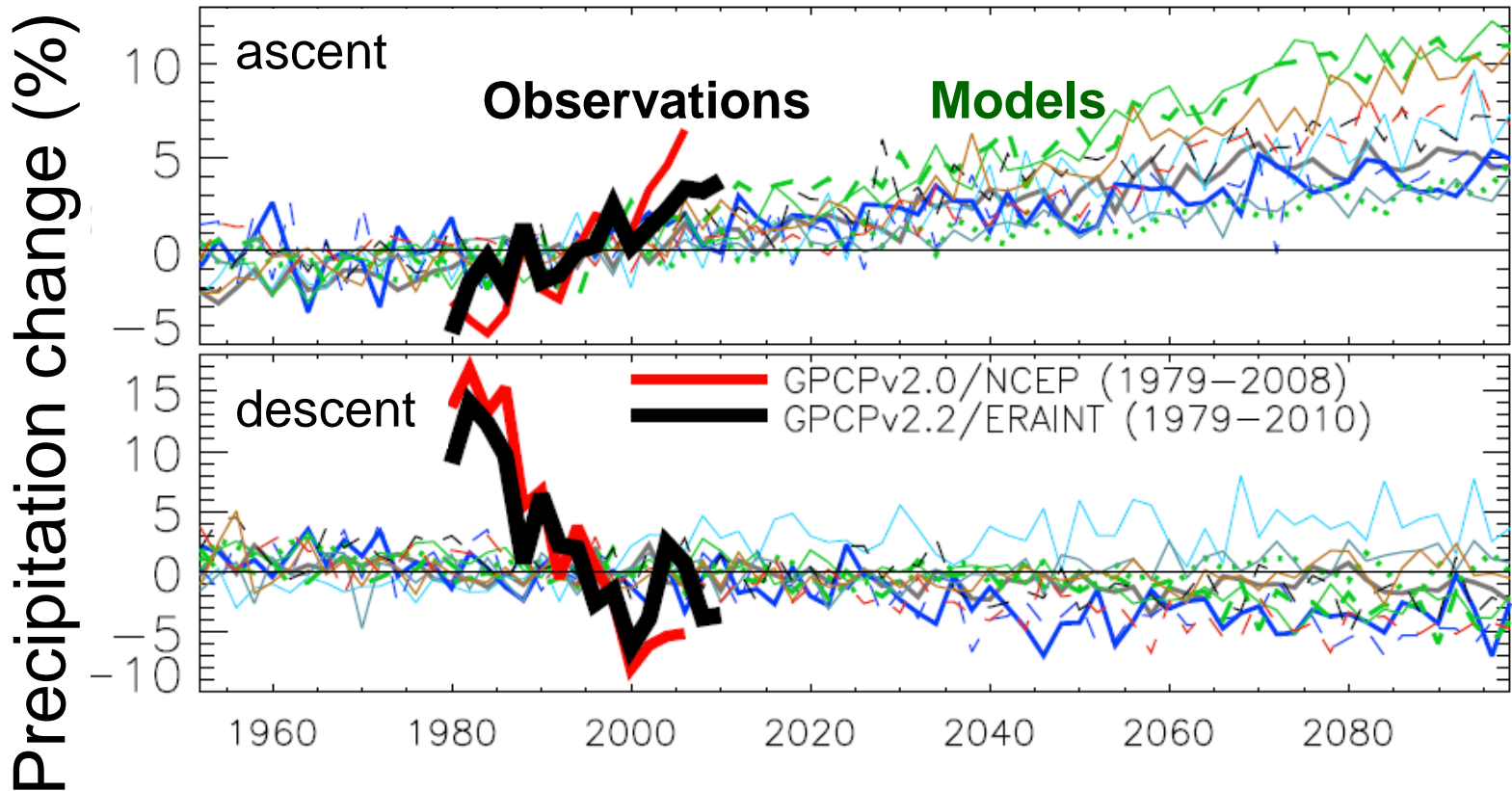
(d) $d\omega_{500}$



Contrasting precipitation response expected



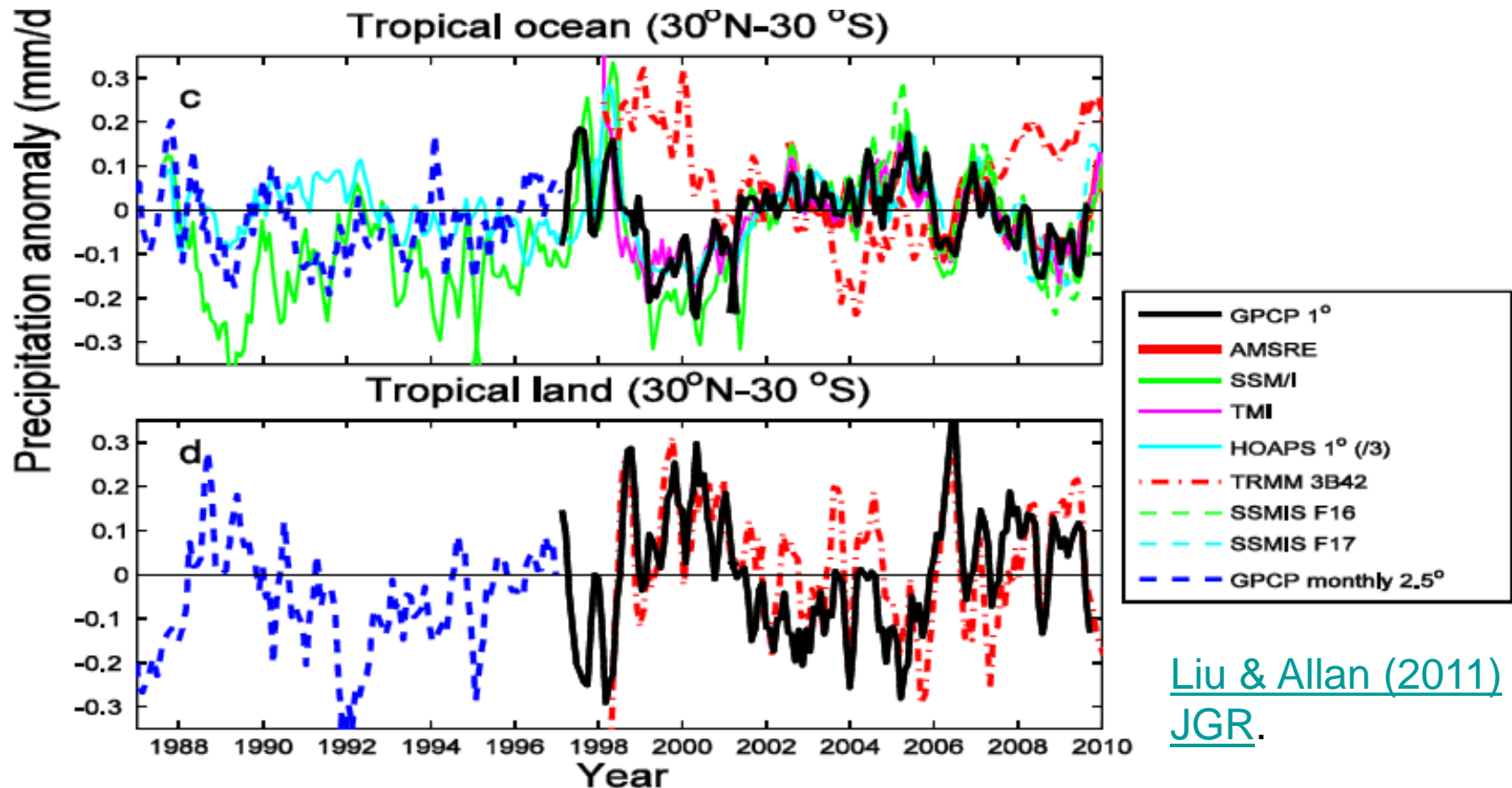
Contrasting precipitation response in wet and dry regions of the tropical circulation



Sensitivity to reanalysis dataset used to define wet/dry regions

Updated from Allan and Soden (2007) GRL; [Allan et al. \(2010\) Environ. Res. Lett.](#)

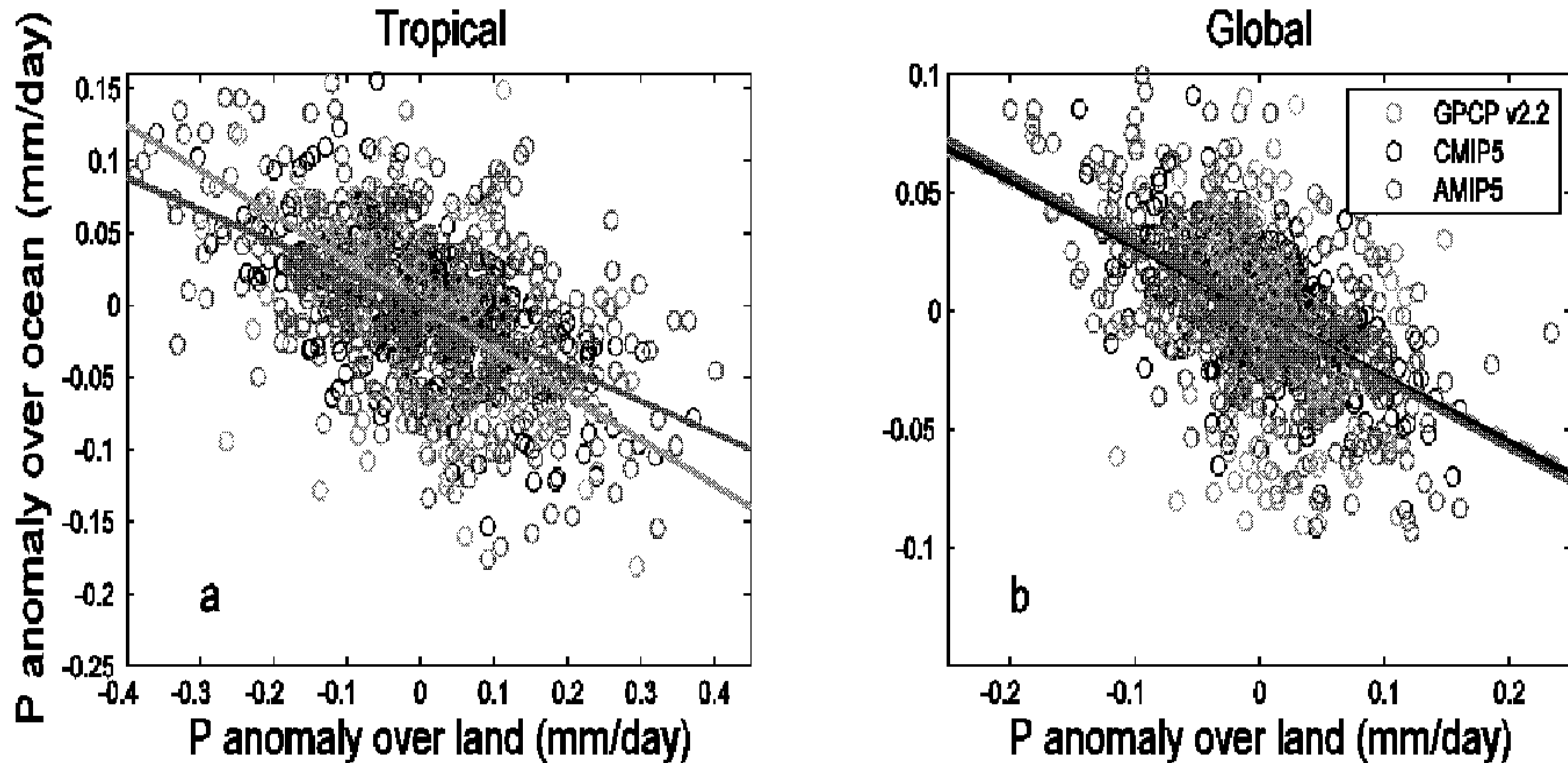
Exploiting satellite estimates of precipitation



[Liu & Allan \(2011\)](#)
[JGR.](#)

- HOAPS and TRMM 3B42 are outliers
- Strong sensitivity to ENSO

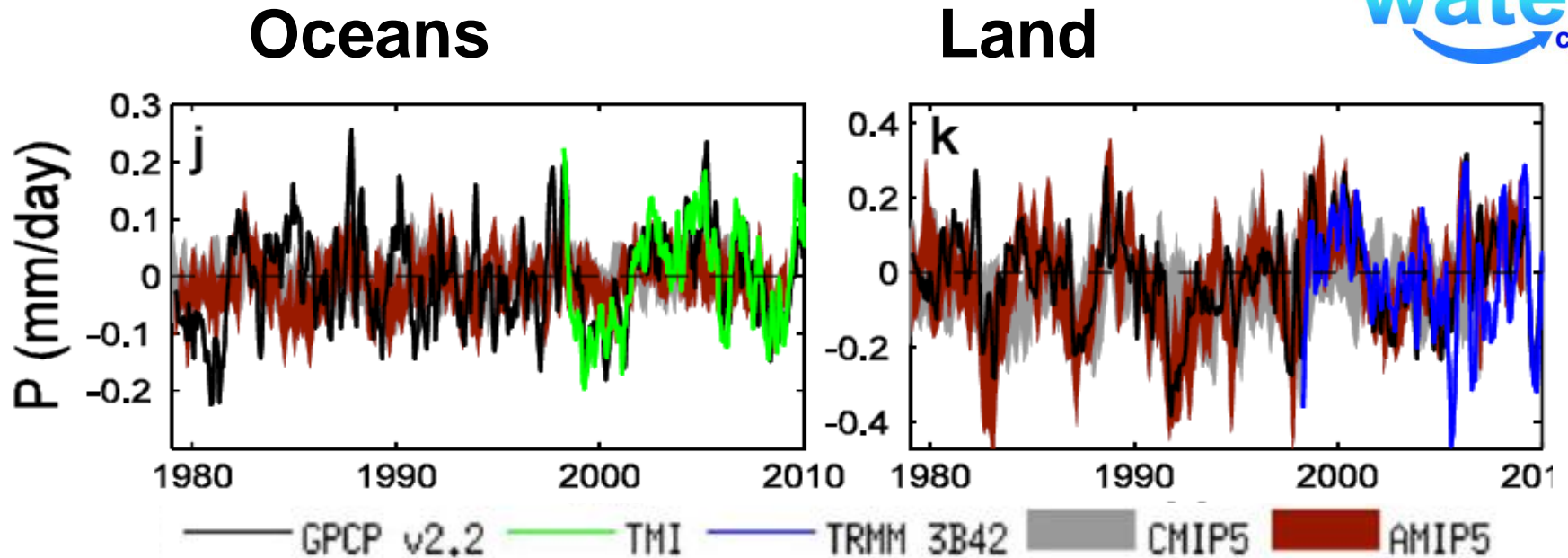
Contrasting land/ocean changes relate to ENSO



Liu and Allan (2012) in preparation

See also Gu et al. (2007) J Clim

PAGODA: Understanding global changes in the water cycle



Above: Current changes in tropical precipitation in CMIP5 models & satellite-based observations

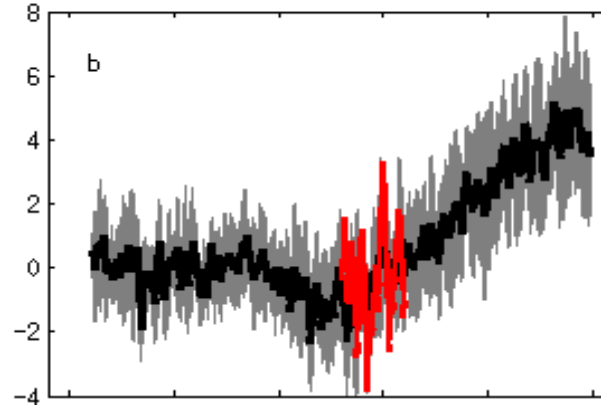
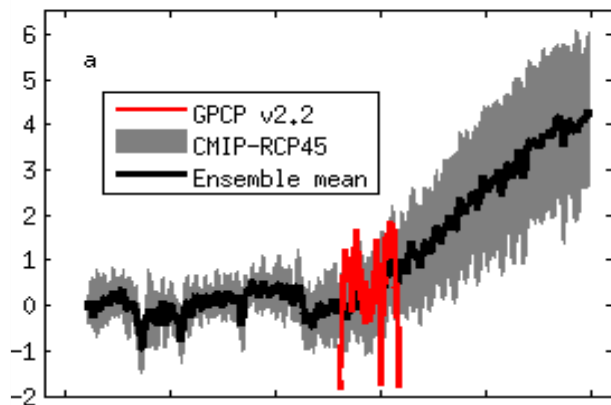
Note realism of atmosphere-only AMIP model simulations

Liu and Allan in prep...

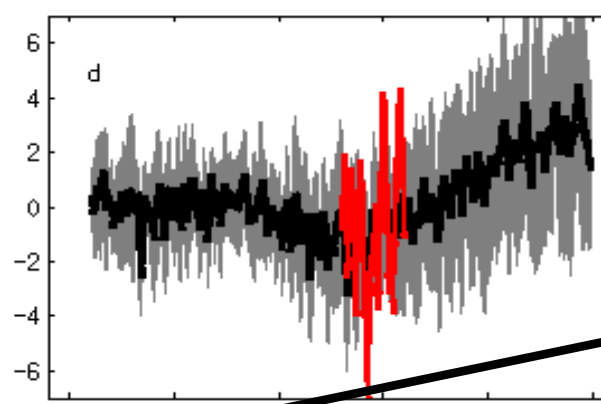
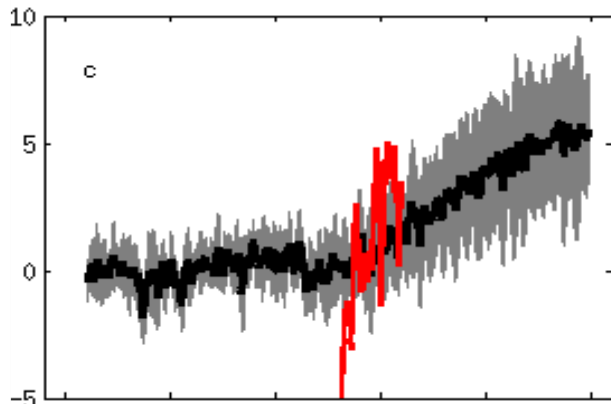
Ocean

Land

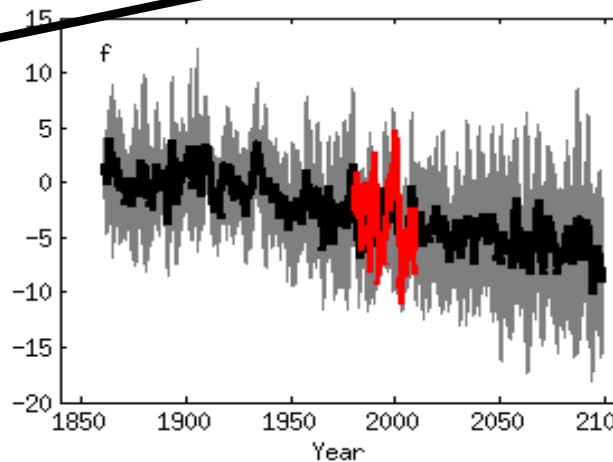
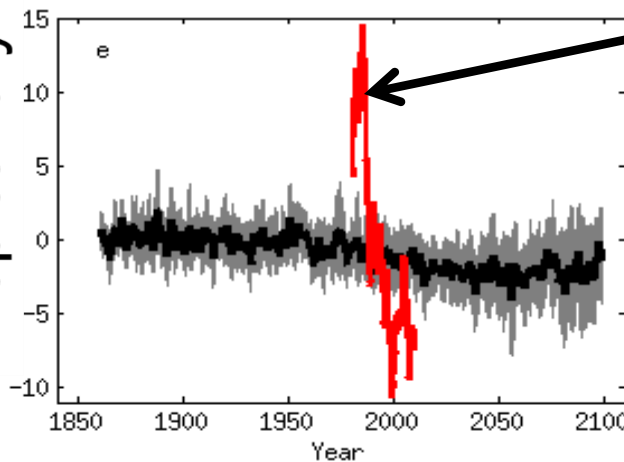
Global



Tropical wet



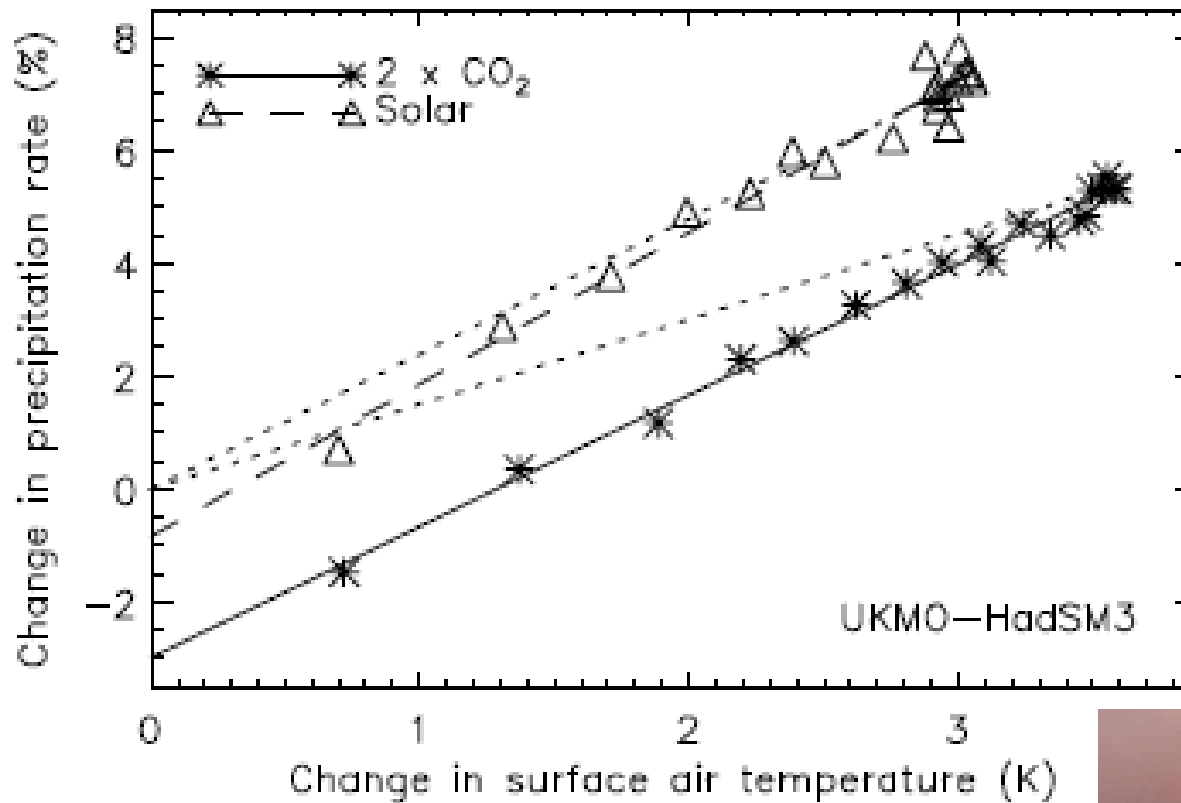
Tropical dry



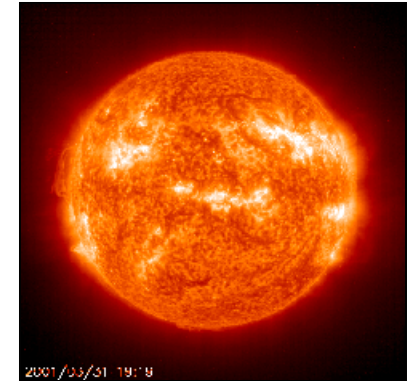
CMIP5
simulated &
projected %
changes in
precipitation

Pre-1988 GPCP
ocean data does
not contain
microwave
observations

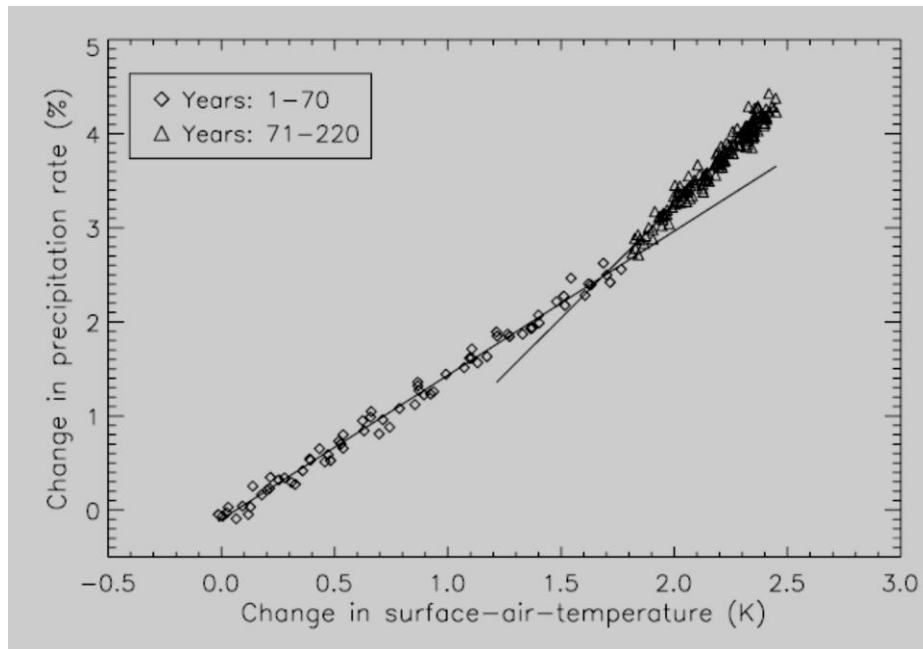
Transient responses



Andrews et al. (2009) J Climate



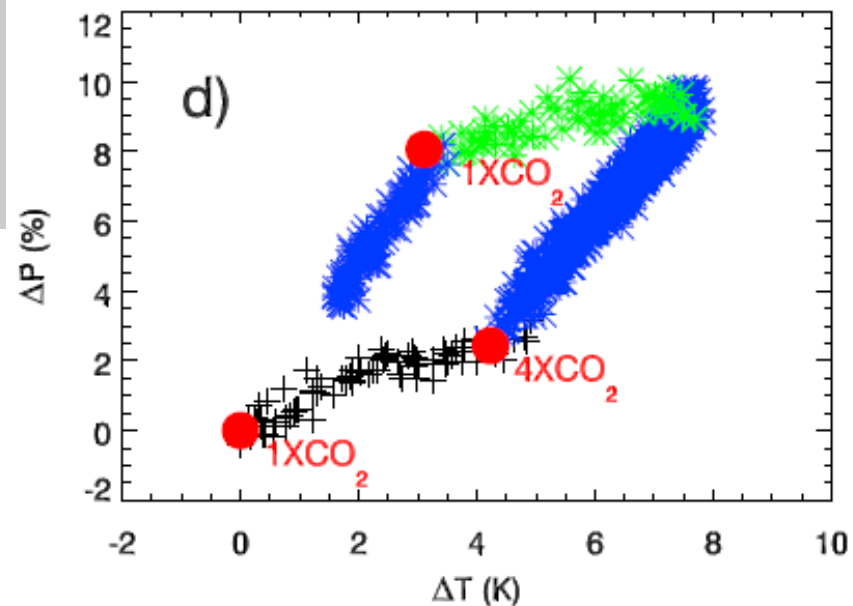
Transient responses



CMIP3 coupled model ensemble mean:
Andrews et al. (2010) Environ. Res. Lett.

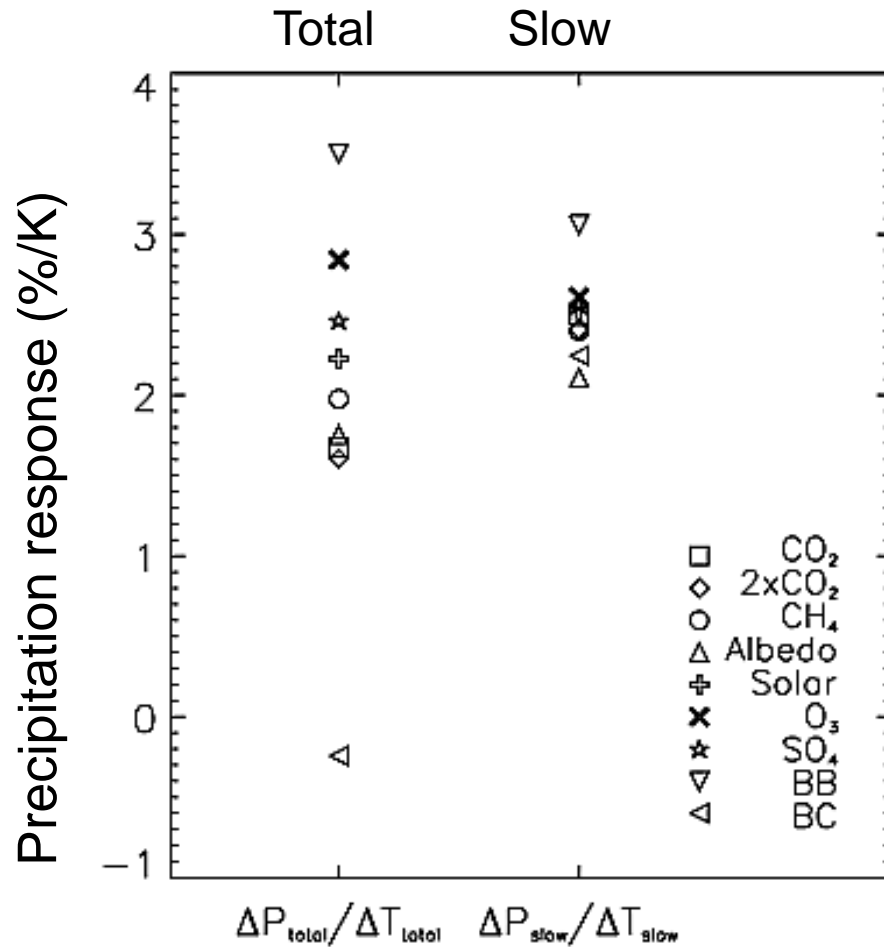
Degree of hysteresis determined by
forcing related fast responses and
linked to ocean heat uptake

- CO₂ forcing experiments
- Initial precip response suppressed by CO₂ forcing
- Stronger response after CO₂ rampdown



HadCM3: Wu et al. (2010) GRL

Forcing related fast responses



Andrews et al. (2010) GRL

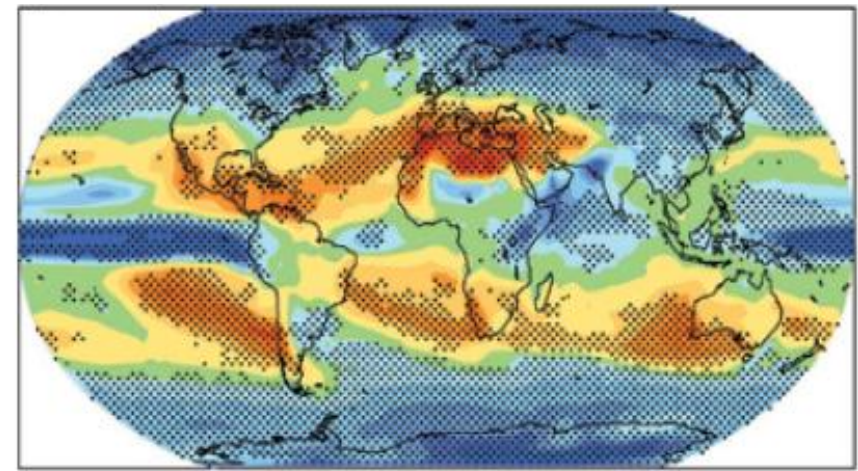
- Surface/Atmospheric forcing determines “fast” precipitation response
- Robust slow response to T
- Mechanisms described in Dong et al. (2009) J. Clim
- CO₂ physiological effect potentially substantial (Andrews et al. 2010 Clim. Dyn.; Dong et al. 2009 J. Clim)
- Hydrological Forcing:
HF=kdT-dAA-dSH

(Ming et al. 2010 GRL; also Andrews et al. 2010 GRL)

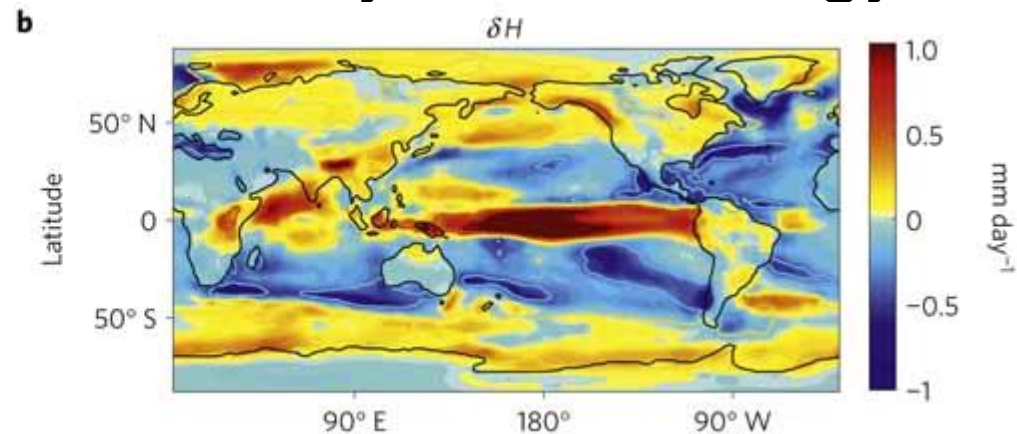
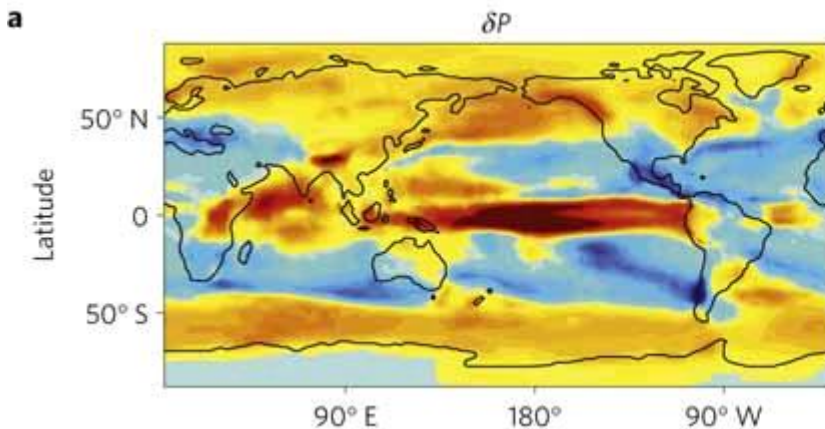
Regional responses in precipitation

Energetic constraints?

- ΔP Precipitation



- ΔH Dry static energy



[Muller and O’Gorman \(2011\)](#) Nature Climate Change
Implications for monsoon? Levermann et al. (2009) PNAS

Conclusions

- **Robust Responses**

- Low level moisture; clear-sky radiation
- Mean and Intense rainfall (roughly)
- Contrasting wet/dry region responses



- **Less Robust/Discrepancies**

- Observed precipitation response at upper end of model range?
- Moisture at upper levels/over land and mean state
- Inaccurate precipitation frequency/intensity distributions
- Magnitude of change in precipitation from satellite datasets/models

- **Further work**

- Decadal changes in global energy budget, aerosol forcing effects and cloud feedbacks: links to water cycle?
- Separating forcing-related fast responses from slow SST response
- Are regional changes in the water cycle, down to catchment scale, predictable?