

Sensitivity of the North Atlantic storm track to regional drivers of change

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CMIP3 Mean response: SRESA1B-20C3M



hPa

20C3M: 19 models (31 ens members), SRESA1B: 19 models (27 ens members)





0.60 0.50 0.40 0:30 0.20 0.10 DJF 0.00 JJA -0.10 -0.20 -0.30 -0.40 -0.50 0.60 20C: 19 models (46 ens members), RCP45: 19 models (30 ens members) hPa

CMIP5 Mean response: RCP45-20C

CMIP3/CMIP5 differences:

- Scenario (SRESA1B vs RCP4.5)
 Ensemble mean Tas responses: CMIP3 = 2.8 K
 CMIP5 = 1.9 K
- CMIP5 higher average resolution
- Several `high-top' models in CMIP5





CMIP5 Mean response: RCP45-20C



CMIP3 Mean response: SRESA1B-20C3M



20C3M: 19 models (31 ens members). SRESA1B: 19 models (27 ens members)

CMIP3: SRESA1B IM SPREAD



hPa

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CMIP3 $\sigma_{IA}/sqrt(n)$



hPa

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CMIP5: RCP45 IM SPREAD



20c: 19 models (46 ens members). rcp45: 19 models (30 ens members)

CMIP5 $\sigma_{IA}/sqrt(n)$



Key question

What are the physical mechanisms causing the inter-model spread in the storm track responses?

Talk Outline

- Part I: What can we learn from the CMIP model runs?
- Part II: Some AGCM experiment results



Climate sensitivity composites

- Based on 14 CMIP5 models
- Scenario: RCP4.5
- Range of Tas responses:

0.7 – 2.8 K

DJF storm track composites based on Global Tas responses



hadgem2 hadgem2-cc miroc-esm canesm2



-0.600.500.400.300.200.100.000.100.200.300.400.500.60 ST response [hPa]

lon

100

Difference

Climate sensitivity composites



DJF storm track composites based on Global Tas responses



hadgem2 hadgem2-cc miroc-esm canesm2



gfdl-esm2g gfdl-esm2m inmcm4 mri-cgcm3

Difference



ion

Meridional T gradient composites





Meridional T gradient composites





Meridional T gradient composites

g

h

2

12

10

8

6

4

2

0

0

Polar Tas response / K



DJF storm track composites based on NH T250 gradients



hadgem2 hadgem2-cc csiro-mk360 ipsl-cm5a-mr



4 models with SMALLEST gradient increase

gfdl-esm2g mri-cgcm3 inmcm4 bcc-csm1-1



Difference



DJF storm track composites based on NH Tas gradients



hadgem2 hadgem2-cc canesm2 miroc-esm

LARGEST

4 models with SMALLEST gradient decrease csiro-mk360 gfdl-esm2g gfdl-esm2m mpi-esm-lr











-0.600.500.400.300.200.100.000.100.200.300.400.500.60 ST response [hPa]

Difference





Consistent with CMIP3/CMIP5 differences?

CMIP3 Mean response: SRESA1B-20C3M



CMIP5 Mean response: RCP45-20C





- Ice losses are similar
- Global Tas response is about 1.5 times larger in CMIP3 than CMIP5





Consistent with CMIP3/CMIP5 differences?

CMIP3 Mean response: SRESA1B-20C3M



CMIP5 Mean response: RCP45-20C



Is this a causal relationship between sea ice and storm tracks?





Conclusions to Part I

- Spread in both CMIP3 and CMIP5 ensembles > sampling error
- Composite plots suggest that meridional temperature gradients can explain some of the spread in the CMIP5 storm track responses
- Southern Hemisphere and Pacific storm tracks: poleward shift is related to change in upper level temperature gradient
- North Atlantic storm track: opposing impacts from upper and lower level temperature gradients
 - Lower level: strong gradient decrease (ice loss?) linked to storm track reduction
 - Upper level: strong gradient increase (large climate sensitivity?) linked to storm track intensification
- However, can't distinguish between local SST changes and larger scale temperature gradients via compositing method

 \rightarrow Motivates experiments



Experimental design

 Force an atmosphere GCM (HadGAM1.2) with a set of SST and sea ice fields designed to capture the spread in the CMIP3 model responses





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Some results...



Styrm risk mitigation



Some results...



a1bp - a1bm



Experimental design

Force an atmosphere GCM (HadGAM1.2) with a set of SST and sea ice fields designed to capture the spread in the CMIP3 model responses



Some results...



Conclusions to Part II

- Atmospheric GCM experiments agree qualitatively with the conclusions from Part I:
 - Global warming with weak ice loss (and cool N Atlantic SSTs) results in a storm track intensification
 - Global warming with strong ice loss (and warm N Atlantic SSTs) results in a storm track weakening
- However, the precise locations of the responses appear to be model dependent
- Split forcing experiments suggest ice loss plays a larger role than sub-polar gyre SSTs in mediating the storm track intensification (although longer runs needed)



What's next?

- Further experiments...
 - A1B- partial-forcing experiments
 - Investigate relationship with upper level temperature gradient (impose a globally uniform SST anomaly?)
 - Repeat experiments using more realistic control SST and sea ice fields
- Further CMIP analysis...
 - Perform compositing/correlation analyses using tracking data

Thank you for listening!

