

# A Basis for Improving Numerical Weather Prediction by Assimilating Doppler Radar Radial Winds

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## What is Data Assimilation (DA)?

What set of Initial Conditions (ICs) will seed the models to best predict the known observations? (**Inverse problem**).

- The **goal of DA** is to **construct** the best possible **ICs**, known as the **analysis**, from which to **integrate** the **NWP** model forward in time.
- **DA** involves deriving the best **current state** of the **atmosphere** by combining a short model forecast (known as **background**<sup>a</sup>) with the latest **observations**, giving each a **weighting** which depends upon their error characteristics.

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<sup>a</sup>**Background** is the most recent information.

## Why use Doppler Radar Data?

### Doppler radar can be used in civilian proposes:

- **Limited Area Models** require **observations** with **high resolution** to generate their **initial conditions** (ICs).
- **Doppler radars**
  - have the ability to **scan large** volumes of atmosphere,
  - provide measurements of **radial velocity** and **reflectivity** with **high** resolution.
- **Doppler winds** give **extra** information in forecasting **quickly** developing **mesoscale systems**.
- The **resolution** of **radar data** are however much **higher** than the **resolution** of the **NWP models**.

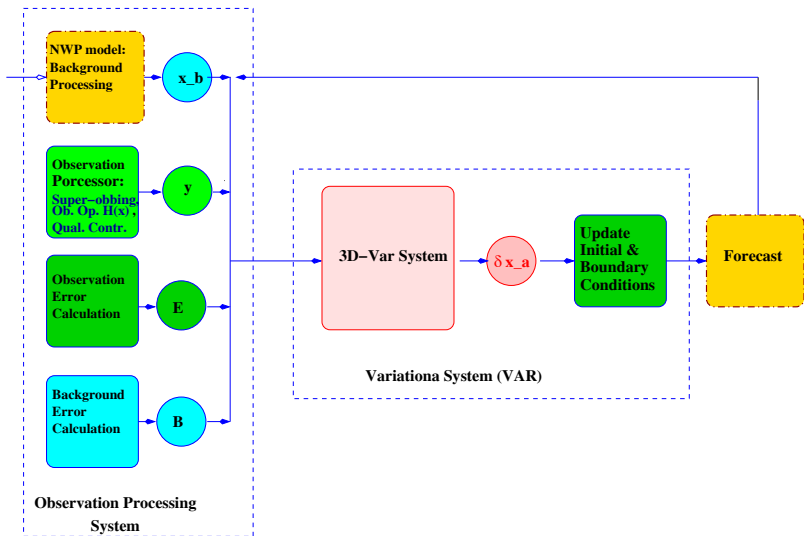
## Steps of NWP

**NWP** is an **IVP**  $\frac{\partial \mathbf{x}}{\partial t} = F(\mathbf{x}, t), \quad \mathbf{x}(t_0) = \mathbf{x}_0$

for which we should provide the **initial conditions** (ICs).

- 1 Collect all possible atmospheric information (**background and observations**) for a given time;
- 2 This information is **analysed** (that represented in **data assimilation**) to **obtain** a **regular, coherent spatial** distribution of the atmosphere at that time;
- 3 This **analysis** becomes the **initial conditions** for the time integration of the **NWP** model;
- 4 Finally, given the estimate of the **present state** of the atmosphere, the model **forecasts** its evolution.

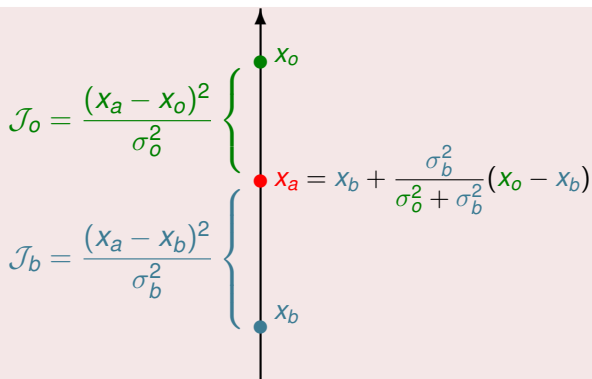
## Continued: NWP Process



## Cost Function

**Total Cost** = **Departure from backg.** + **Departure from obs.**

$$\mathcal{J}[x_a] = \mathcal{J}_b[x_a] + \mathcal{J}_o[x_a]$$



## Continued: Cost Function

$$\mathcal{J}[\mathbf{x}] = \frac{1}{2}(\mathbf{x}_b - \mathbf{x})^T \mathbf{B}^{-1}(\mathbf{x}_b - \mathbf{x}) + \frac{1}{2}(\mathbf{y} - \mathcal{H}[\mathbf{x}])^T \mathbf{E}^{-1}(\mathbf{y} - \mathcal{H}[\mathbf{x}]).$$

$\mathbf{x}$  denotes the analysis,  $\mathbf{x}_b$  background,  $\mathbf{y}$  observation vector.  $\mathbf{B}$  and  $\mathbf{E}$  are back. and obs. error covariance matrices.

$\mathcal{H}$  is the observation operator that relates the model variables to the observation variable and a transformation between the different grid meshes.

$\mathcal{H}$  can be linearized as  $\mathcal{H}[\mathbf{x} + \delta\mathbf{x}] = \mathcal{H}[\mathbf{x}] + \mathbf{H}\delta\mathbf{x}$ . Then

$$\mathcal{J}[\delta\mathbf{x}] = \frac{1}{2}\delta\mathbf{x}^T \mathbf{B}^{-1}\delta\mathbf{x} + \frac{1}{2}[\mathbf{H}\delta\mathbf{x} - \mathbf{y} + \mathcal{H}\mathbf{x}_b]^T \mathbf{E}^{-1}[\mathbf{H}\delta\mathbf{x} - \mathbf{y} + \mathcal{H}\mathbf{x}_b],$$

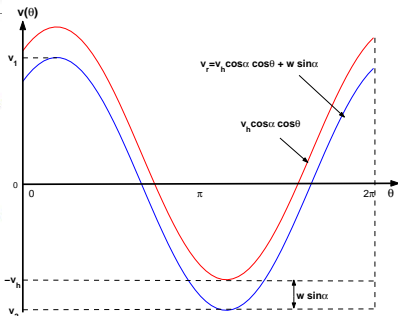
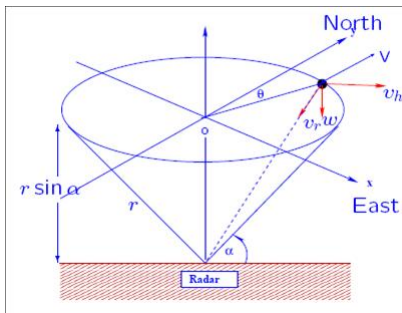
where  $\mathbf{H}$  is linear operator  $\equiv \partial\mathcal{H}/\partial\mathbf{x}$ .



## Objectives and Challenges

- 1 The **aim** is to **develop** an incremental assimilation of **Doppler radar winds** using **3D-Var/4D-Var**, leading to an **improved representation** of the **Initial Conditions (ICs)** for **storm-scale forecasting of winds, sand or dust in UAE!!**.
- 2 **Main Challenges:**
  - Data assimilation at **high resolution**; loss of **vertical resolution** of the radar data.
  - The observed quantities are **not model variables**.
- 3 **Organization:**
  - Obtain data for **radial winds** for which the **model output** is available,
  - The very high resolution raw data is **re-mapped** to model resolution, and make it **suitable** for input to the **Var** system,
  - Formulate an **observation operator**,
  - **Assimilate** the processed data, and investigate the **impacts** on the **analysis** and **model forecasts**.

## Pre-processing the Doppler radar data

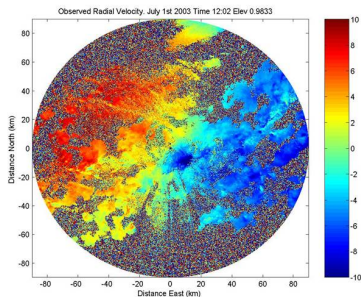
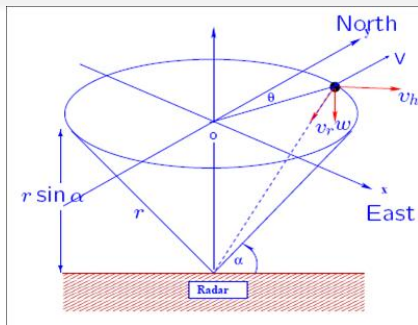


- Bilinear interpolation of  $u$ ,  $v$  &  $w$  to the obs locations.
- Projection of the horizontal model wind towards the radar beam  
 $v_h = u \sin \theta + v \cos \theta$ , where  $\theta$  is the azimuth angle.
- The radial wind  $v_r$  is  $v_h$  + the vertical velocity term, in direction of the radar beam,

$$v_r = u \cos \alpha \sin \theta + v \cos \alpha \cos \theta + w \sin \alpha.$$

$\alpha$  is the elevation.

## Getting the Doppler radar data

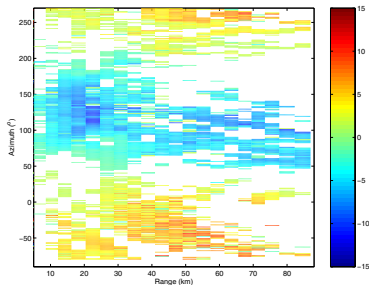
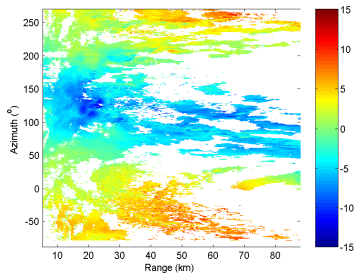
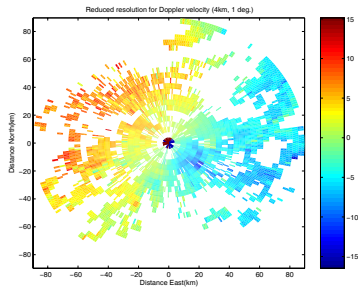
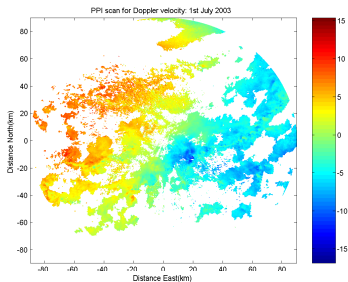


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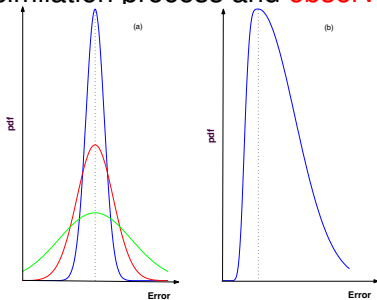
$\alpha$  is the elevation.

# Pre-processing data - Super-obing



## Types of Errors in Doppler Radial winds

- **Errors in the original measurements** with each pulse volume that **depends** on the **velocity gradients** within the pules volume (vary with range);
- **Errors due to super-obbing** procedured (vary with range);
- **Errors due to hardware** degradation (random);
- **Errors** in the assimilation process and **observation operator**.



Distribution of wind speed error due to (a) **instrumental noise**, (b) **strong velocity gradient across the pulse volume**.

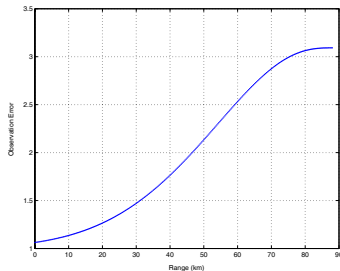
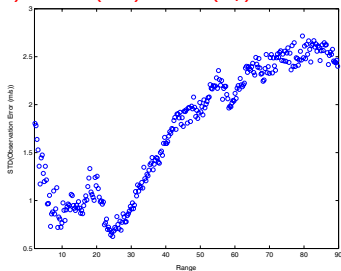
## Continued: Types of Errors in Doppler Radial winds

- The **obs errors** are **uncorrelated** in space and time.
- The **error in radial winds** due to the velocity gradient along the pulse volume, **varies** with the range  $R$ , is

$$\check{\sigma}^2(\varepsilon_v) = \left(1 - e^{-|\Delta v_r/v_r|}\right) \sigma^2(v_r).$$

- The **instrumental error**  $\hat{\sigma}^2(\varepsilon_i)$  does not vary temporally.
- The **total error variance** of the radial winds is

$$\sigma^2(\varepsilon) = \check{\sigma}^2(\varepsilon_v) + \hat{\sigma}^2(\varepsilon_i).$$



Local **Observation errors** (m/sec) of the radial velocity as a function of the **range** (left), proposed s-function (right).

## 3D-Var System

- The Met Office **3D-Var** system uses the incremental cost function,

$$\mathcal{J}[\delta\mathbf{x}] = \frac{1}{2}\delta\mathbf{x}^T \mathbf{B}^{-1}\delta\mathbf{x} + \frac{1}{2}[\mathbf{H}\delta\mathbf{x} - \mathbf{y} + \mathcal{H}\mathbf{x}_b]^T \mathbf{E}^{-1}[\mathbf{H}\delta\mathbf{x} - \mathbf{y} + \mathcal{H}\mathbf{x}_b],$$

$\mathbf{H}$  is linear operator  $\equiv \partial\mathcal{H}/\partial\mathbf{x}$ .

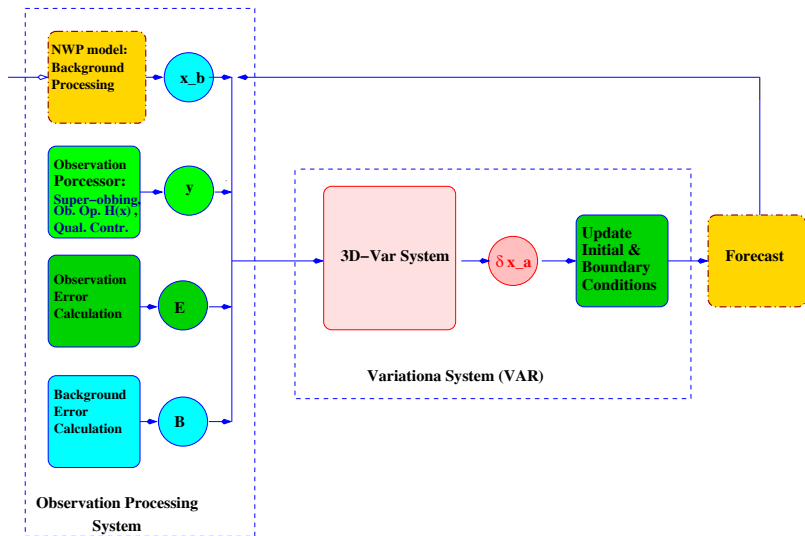
- To avoid inversion of  $\mathbf{B}$ , we use ( $\mathcal{X} = \mathbf{U}\delta\mathbf{x}$ )

$$\mathcal{J}[\mathcal{X}] = \frac{1}{2}\mathcal{X}^T \mathcal{X} + \frac{1}{2}[\mathbf{H}\mathbf{U}^{-1}\mathcal{X} - \mathbf{y} + \mathcal{H}\mathbf{x}_b]^T \mathbf{E}^{-1}[\mathbf{H}\mathbf{U}^{-1}\mathcal{X} - \mathbf{y} + \mathcal{H}\mathbf{x}_b].$$

$\mathbf{U}$  transforms the **forecast error** in the physical model space into a **space** where the covariance is **identity** matrix.

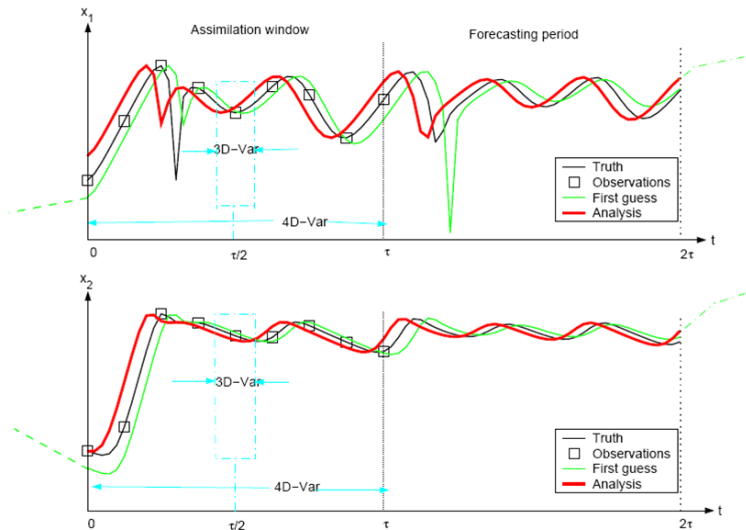
- $\mathbf{E}$  is a diagonal matrix, includes: **observation errors** + error from **super-obbing** + errors from the **observation operator**.

## Continued: 3D-Var system



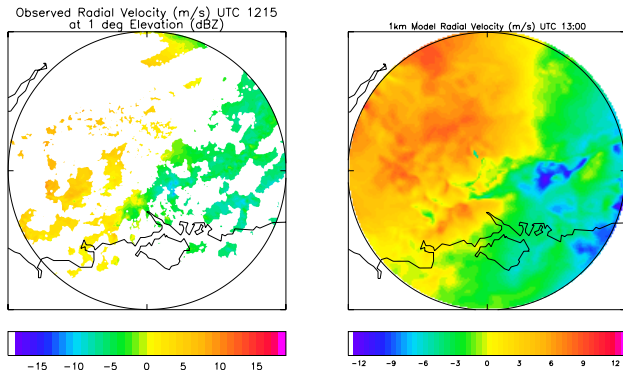


## Continued: 3D-Var system



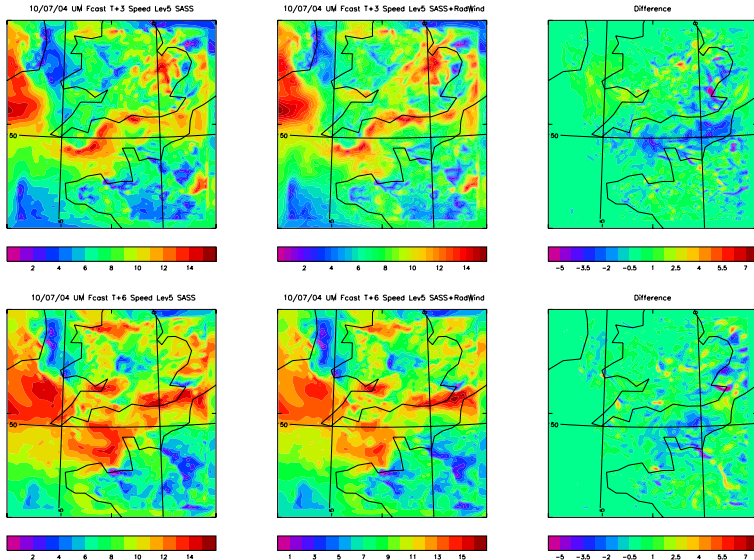
## Case Study: Assimilating Doppler Radial Velocity

Observed radial winds for the case study in UK have been assimilated using 3D-Var system in the Met Office. We insert total errors varying with range from 2m/s to 6m/s.



Observed radial velocity compared with assimilated radial velocity, when running 3D-Var system at 4km resolution.

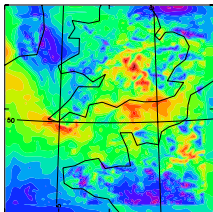
# UM Forecasts 10<sup>th</sup> July 2004 Case



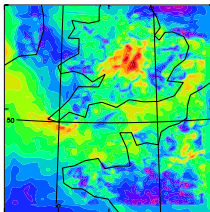
Wind speeds at  $T + 3$  (top) &  $T + 6$  (bottom), level 5 from 12UTC. From left to right: radial wind, and the difference.

## Continued: UM Forecasts

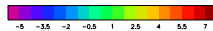
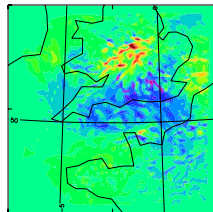
10/07/04 UM Fcast T+1 Speed Lev5 SASS



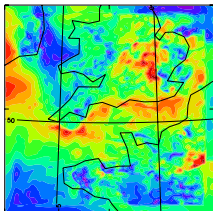
10/07/04 UM Fcast T+1 Speed Lev5 SASS+RadWind



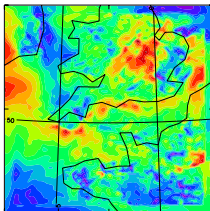
Difference



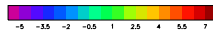
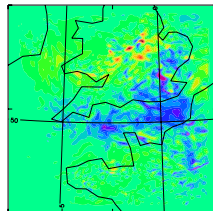
10/07/04 UM Fcast T+2 Speed Lev5 SASS



10/07/04 UM Fcast T+2 Speed Lev5 SASS+RadWind



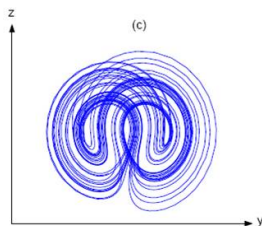
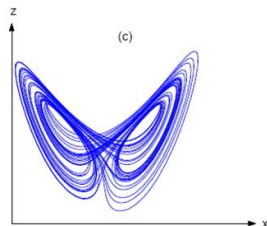
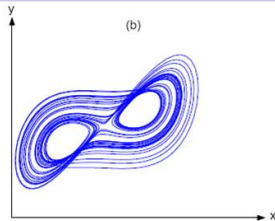
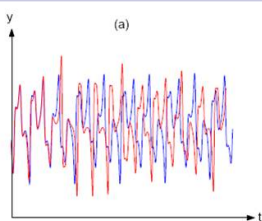
Difference



## Why Are Weather Forecasts Sometimes Wrong?

There are two main reasons: (i) Sensitivity & (ii) Nonlinearity.

### (i) Sensitivity to the ICs



## Continued

## (ii) Biases due to nonlinearity

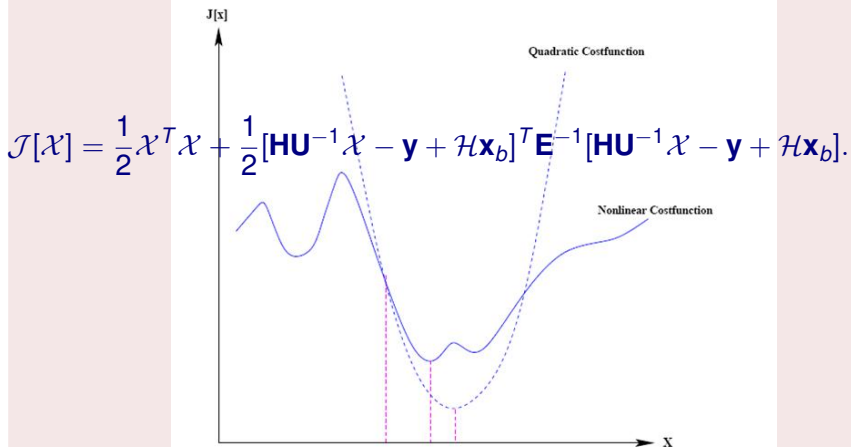


Fig. 4. Iterative solution for quadratic and nonlinear cost function.

## Summary & Future Directions

### Steps & Requirements:

- **Observation processing and data collections;**
- **Variation systems and analysis;**
- **Forecast.**






### Future Directions:

- **Assimilate radial velocity, using 4D-Var technique to produce ICs for 1-4km resolution forecasts.**
- **Applications to sand and dust storm in Abu Dhabi Area.**

### Acknowledgment:

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## References:

-  F.A. Rihan, M. Anwar, C. Collier, **A Basis for Improving Numerical Forecasting in the Gulf Area by Assimilating Doppler Radar Radial Winds**, *Int. J. Geosc.*, (2010), 70–78.
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-  F.A. Rihan, C.G. Collier, S.P. Ballard, **Impact of assimilation of Doppler radial velocity on a variational system and on its forecasts**, *AMS*, October (2005) 24-29.
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