5th International Symposium on Data Assimilation, July 18–22, 2016, Reading, U.K.

Ensemble-based Experimental Atmospheric Reanalysis Using a Global Coupled Atmosphere–Ocean GCM



Nobumasa Komori,¹ Takeshi Enomoto,^{1,2} Takemasa Miyoshi,^{1,3,4} Akira Yamazaki,¹ Akira Kuwano-Yoshida,¹ and Bunmei Taguchi^{1†}

¹ Application Laboratory, Japan Agency for Marine-Earth Science and Technology, Yokohama, Japan

- ² Disaster Prevention Research Institute, Kyoto University, Uji, Kyoto, Japan
- ³ RIKEN Advanced Institute for Computational Science, Kobe, Japan
- ⁴ Department of Atmospheric and Oceanic Science, University of Maryland, College Park, MD, U.S.A.
- ⁺ Now at Research Center for Advanced Science and Technology, University of Tokyo, Tokyo, Japan

É-mail: komori@jamstec.go.jp; http://www.jamstec.go.jp/res/ress/komori/

To enhance the capability of the local ensemble transform Kalman filter (LETKF) with the Atmospheric general circulation model (GCM) for the Earth Simulator (AFES), a new system has been developed by replacing AFES with the Coupled atmosphere-ocean GCM for the Earth Simulator (CFES). An initial test of the prototype of the CFES–LETKF system has been completed successfully, assimilating atmospheric observational data (NCEP PREPBUFR archived at UCAR) every 6 hours to update the atmospheric variables, whereas the oceanic variables are kept unchanged throughout the assimilation procedure.

An experimental retrospective analysis–forecast cycle with the coupled system (CLERA-A) starts

1. Introduction

on-the-fly estimation of analysis and forecast errors, 2013]. relative ease of implementation, and efficiency with parallel computers.

(ALEDAS2), assimilating observational data of the Na- similation system (CLEDAS-A).

Ensemble-based data assimilation techniques have been tional Centers for Environmental Prediction (NCEP) global rapidly growing because of their advantages of the data assimilation system (PREPBUFR) [Enomoto et al.,

In ensemble data assimilation systems based on atmospheric GCMs (including ALEDAS), however, surface Miyoshi and Yamane [2007] applied the local en- boundary conditions such as sea surface temperature semble transform Kalman filter (LETKF) to an atmo- (SST) and sea-ice distribution are the same among all spheric general circulation model (GCM), AFES, to con- ensemble members, which leads to an underestimation struct the AFES-LETKF ensemble data assimilation of the ensemble spread (an overestimation of the foresystem (ALEDAS). Miyoshi et al. [2007] performed one cast accuracy) near the surface. Additionally air-sea and a half years of AFES–LETKF experimental ensemble coupled phenomena, e.g., lead–lag relationship between reanalysis (ALERA) using observational dataset of the SST and precipitation over the tropics, are not well repro-Japan Meteorological Agency operational system. Cur- duced in such systems. To overcome these problems, we rently the second generation of ALERA (ALERA2) is un- replace AFES with a coupled atmosphere-ocean GCM, derway with the latest version of AFES and LETKF **CFES**, to develop the CFES-LETKF ensemble data as-

on August 1, 2008, and the atmospheric initial conditions (63 members) are taken from the second generation of AFES-LETKF experimental ensemble reanalysis (ALERA2). The ALERA2 analyses are also used as forcing of stand-alone 63-member ensemble simulations with the Ocean GCM for the Earth Simulator (EnOFES), from which the oceanic initial conditions for the CLERA-A are taken.

The ensemble spread of SST is larger in CLERA-A than in EnOFES, suggesting positive feedback between the ocean and the atmosphere. Although SST in CLERA-A suffers from the common biases among many coupled GCMs, the ensemble spreads of air temperature and specific humidity in the lower troposphere are larger in CLERA-A than in ALERA2. Thus replacement of AFES with CFES successfully contributes to mitigate an underestimation of the ensemble spread near the surface resulting from the single boundary condition for all ensemble members and the lack of atmosphere-ocean interaction.

In addition, the basin-scale structure of surface atmospheric variables over the tropical Pacific is well reconstructed from the ensemble correlation in CLERA-A but not ALERA2. This suggests that use of a coupled GCM rather than an atmospheric GCM could be important even for atmospheric reanalysis with an ensemble-based data assimilation system.

3. Results

3.1. Comparison of ALERA2 and CLERA-A







2. CFES–LETKF Ensemble Data Assimilation System

The resolution of the atmospheric component of CFES used in the system is T119 (~100 km) in the horizontal and **48 layers** in the vertical, the same as in ALERA2. The oceanic component has a resolution of 1/2° (~50 km) in the horizontal and 54 levels in the vertical, and is coupled with the atmospheric component every hour.

An experimental retrospective analysis-forecast cycle with CLEDAS-A (63 members) is conducted from August 1 to September 30, 2008, assimilating the atmospheric observational data (NCEP PREPBUFR archived at UCAR) every 6 hours to update the atmospheric variables, whereas the oceanic variables are kept unchanged throughout the assimilation procedure. The result is referred to as **CLERA-A**.

The 63-member initial conditions of the atmospheric component are taken from ALERA2. The initial conditions of the oceanic component are made by stand-alone ensemble ocean simulations (EnOFES) forced by each member of ALERA2 after 60-year spin-up integration.



Figure 1: The data flow charts of an analysis–forecast cycle for (a) ALEDAS2 and (b) CLEDAS-A. Rectangles represent input/output data, and round rectangles represent processes.

3.2. Ensemble Covariance and Correlation

Ens. Covar. b/w Surf. Temp. & Cumulus Precip. (Lag: +1d)



Ens. Covar. b/w Surf. Temp. & Cumulus Precip. (Lag: ±0d)



Ens. Covar. b/w Surf. Temp. & Cumulus Precip. (Lag: -1d)



(d) Ens. Covar. b/w Surf. Temp. & Large-Scale Precip. (Lag: +1d)



(e) Ens. Covar. b/w Surf. Temp. & Large-Scale Precip. (Lag: ±0d)



(f) Ens. Covar. b/w Surf. Temp. & Large-Scale Precip. (Lag: -1d)



Figure 5: Lag ensemble covariance between forecasted SST and precipitation induced by (left) cumulus convection and (right) large-scale condensation in CLERA-A. (Top) SST leads precipitation 1 day, (middle) simultaneous, and (bottom) precipitation leads SST 1 day.



Figure 4: Ensemble spread of (left) forecasted air temperature and (right) specific humidity at the height of $\sigma = 0.97$ averaged over the experimental period for (top) ALERA2, (middle) CLERA-A, and (bottom) their difference (CLERA-A minus ALERA2).



Figure 6: Vertical ensemble correlation between forecasted SST and air temperature at the height of (a) $\sigma = 0.51$ and (b) $\sigma = 0.86$ in CLERA-A averaged over the experimental period.

Figure 7: Ensemble correlation of forecasted 2-m air temperature (color), 10-m winds (green arrows), and surface pressure (gray contour) with respect to 2-m air temperature in the NINO3.4 region (indicated by a white rectangle) averaged over the experimental period for (a) ALERA2 and (b) CLERA-A.